

Impact of nonlinear energy transfer on the wave field

Tamura et al. (2010): JGR -Oceans

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Motivations

The 3rd generation wave model

$$\frac{\partial F}{\partial t} + c_g \nabla F = \underbrace{S_{in} + S_{ds}}_{\text{external source: } Sf} + S_{nl}$$

resonant interaction

S_{nl} controls the evolution of wave spectra

1. To understand whether accurate ***S_{nl}*** scheme improve the model representation of wave parameters and spectral shapes.
2. To investigate the role of ***S_{nl}*** in the source balance in conjunction with the parameterization of the external source ***S_f***.

Methodologies

SnI schemes:

DIA:	Hasselmann et al. (1985)
TSA:	Resio and Perrie (2008)
Multiple DIA or XDIA:	Van Vledder (2001), Tolman (2004)....
SRIAM method:	Komatsu (1996), Tamura et al. (2008)
RIAM:	Komatsu and Masuda (1996)
WRT:	Van Vledder (2006)
Masuda method:	Masuda (1980)
...	

Pacific hindcast experiments:

WAVEWATCH-III v2.22 (Tolman, 2002)

Wind input & Dissipation : Tolman & Chalikov (1996)

Validation:

NOAA/NDBC buoys –Hs, Tp, Freq. spectra-
(46035, 46066, 46005, 46006, 46089, 51001, 51004, 51028)

Conclusions

- A negligible difference between SRIAM and DIA for H_s . However, **the difference for the T_p was quite pronounced, especially around the tropical Pacific**, where a persistent bias in T_p was improved by using SRIAM.
- SRIAM can quantitatively capture the overshoot phenomena around the spectral peak during wave growth.
- S_{nl} played a major role in maintaining the equilibrium range; **it reacted to changes in the net external sources to cancel out the total source term.**
- The magnitude of external source controls the spectral tail exponent **in the equilibrium range so as to support Resio et al. (2004).**

Numerical treatment of the nonlinear transfer function

SRIAM method: Komatsu (1996), Tamura et al. (2008)

an efficient scheme for operational use

DIA : 1 resonant configuration

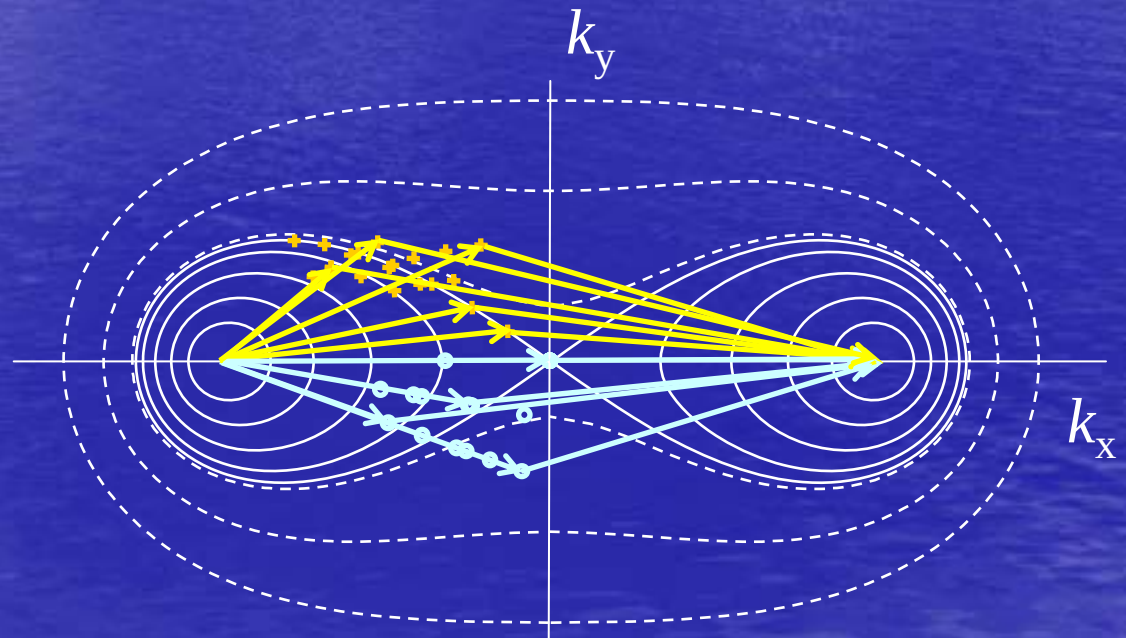
SRIAM : 20 resonant configurations

Exact method: 10^3 to 10^4 resonant configurations

Resonance condition

$$\mathbf{k}_1 + \mathbf{k}_2 = \mathbf{k}_3 + \mathbf{k}_4$$

$$\sigma_1 + \sigma_2 = \sigma_3 + \sigma_4$$



20 optimized resonant configurations

2004 hindcast experiments

Model configuration

Third-generation wind-wave model

WAVEWATCH-III v2.22 (Tolman, 2002)

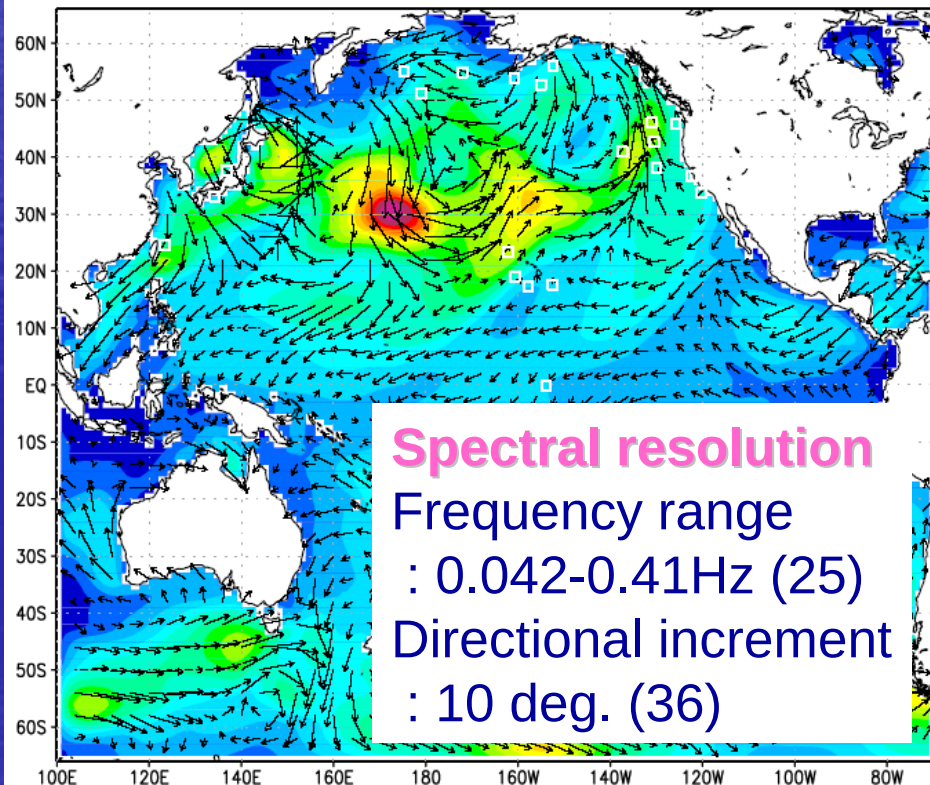
Wind input & Dissipation
: Tolman & Chalikov (1996)

Nonlinear transfer function
: SRIAM method

non-parametric spectral tail

Surface wind field:
NCEP/NCAR Reanalysis

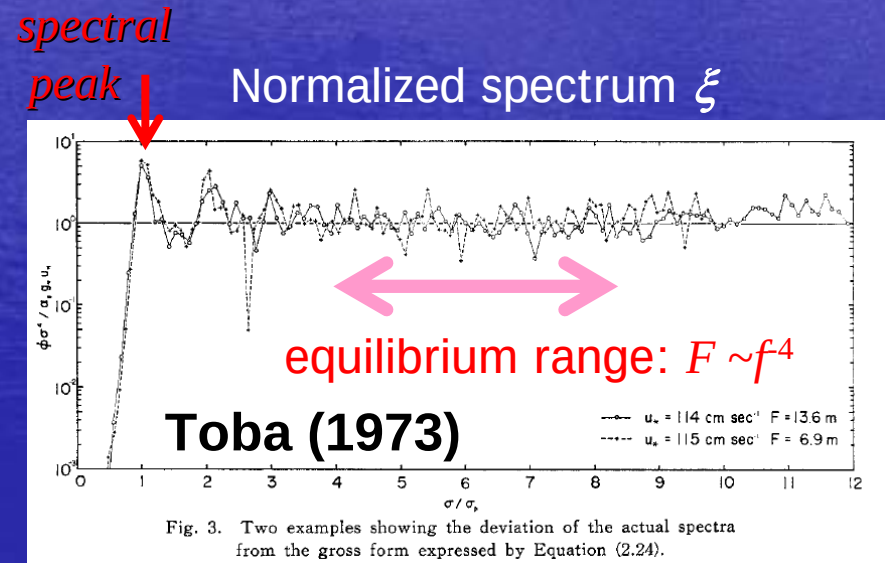
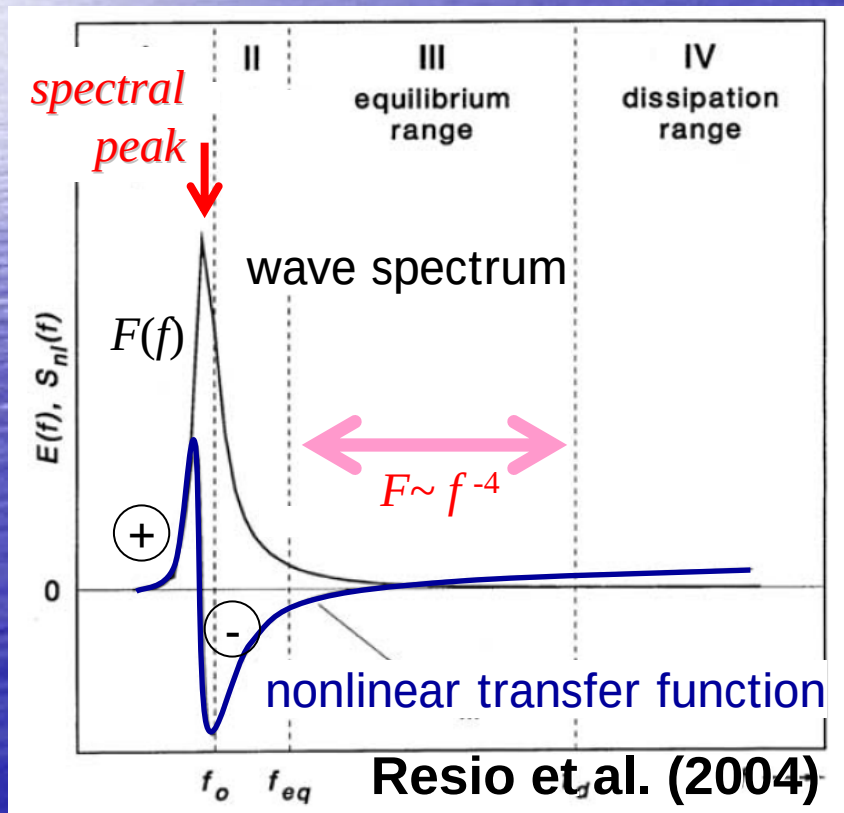
Pacific model ($1^\circ \times 1^\circ$)



Evolution of wave spectra in the equilibrium range

$$\frac{\partial F}{\partial t} = S_{in} + S_{nl} + S_{ds} \approx 0$$

To maintain the equilibrium range, the sum of the three source terms should be zero within the equilibrium range



$$\xi = \frac{F(f) f^4}{\alpha \cdot u_* g}$$

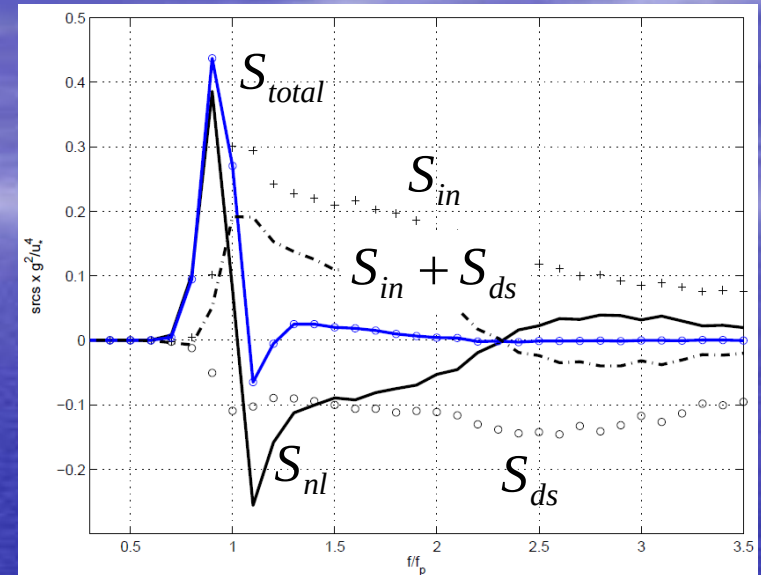
Wave spectral shape and source balance

S_{in} : Tolman & Chalikov (1996)

$$S_{ds} = S_{ds}^{low} + \beta \cdot S_{ds}^{high}$$

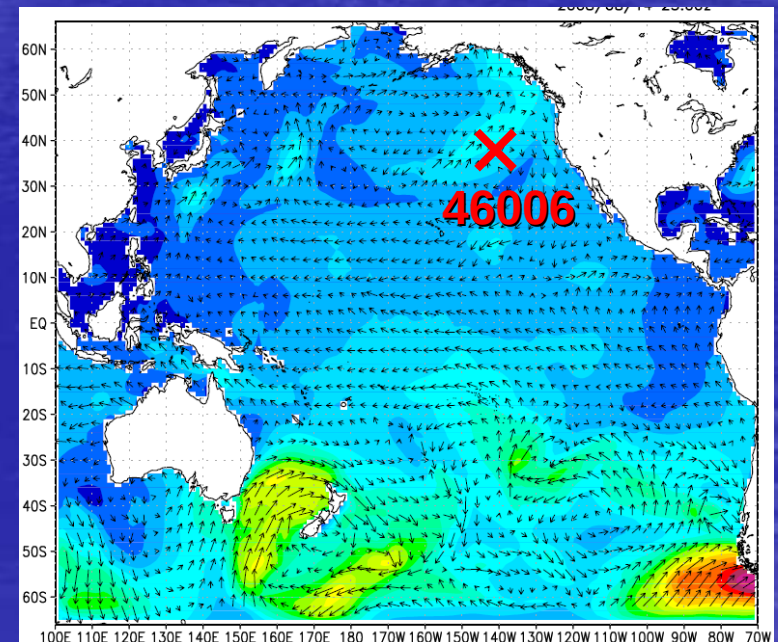
$\beta = 0, 0.25, 0.5, 0.6, 0.75, 1$

S_{nl} : SRIAM Komatsu (1996)

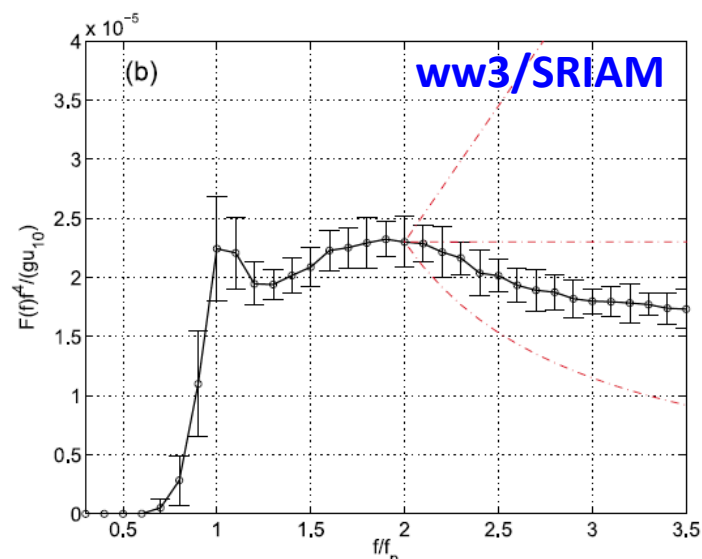
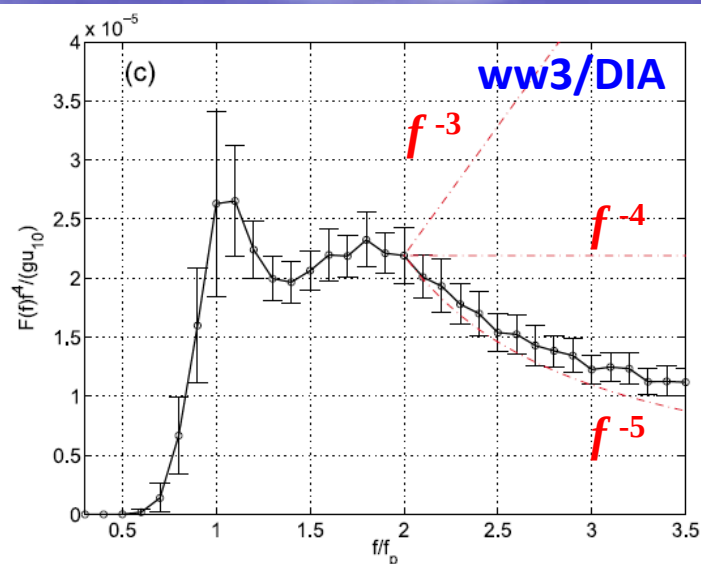
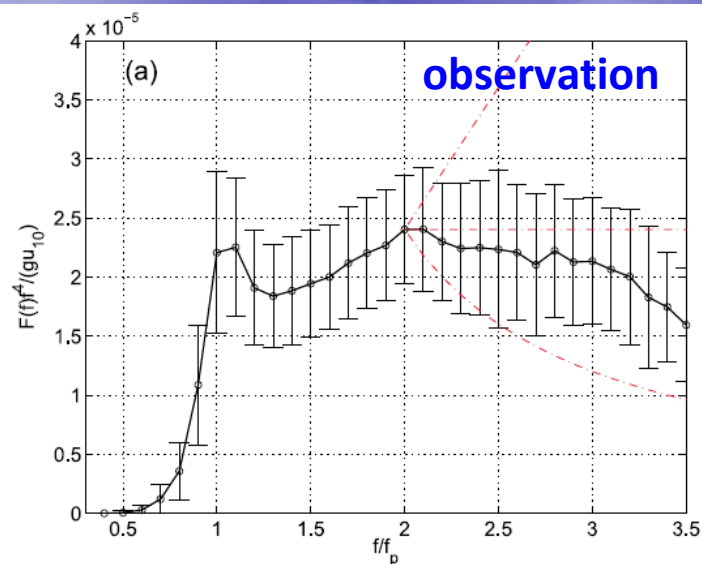


Wave spectra investigated

1. single peak
2. for growing sea state ($U_{10}/c_p > 1$)



Mean wave spectra @46006



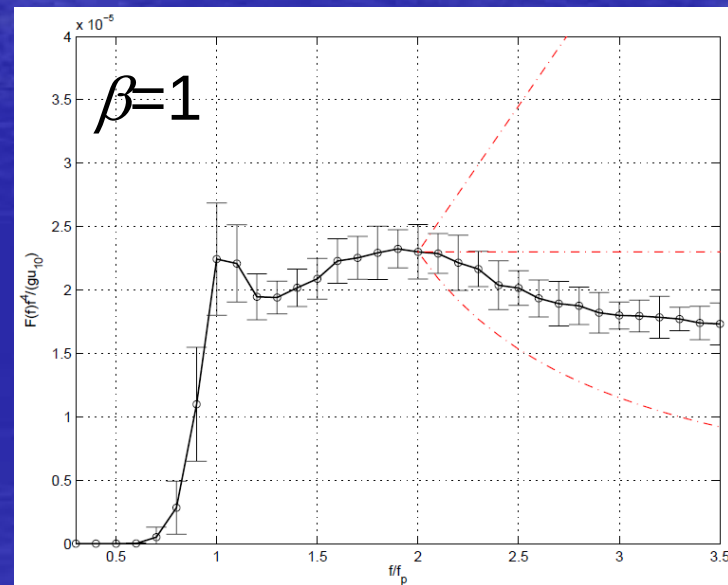
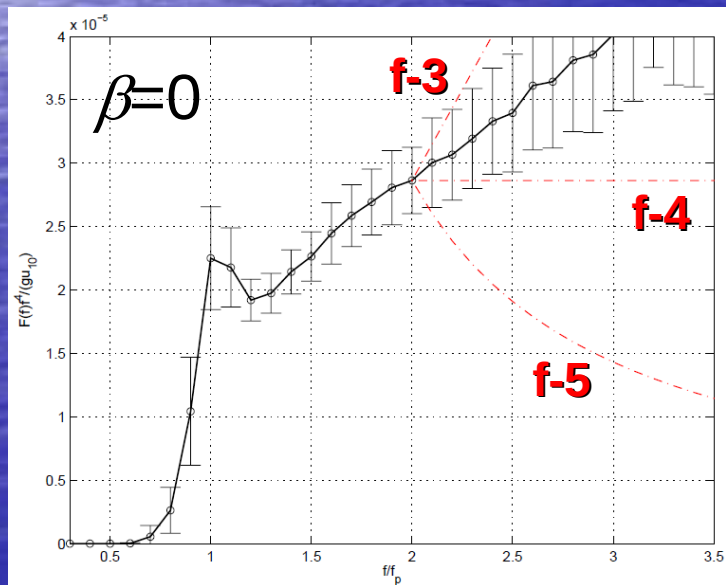
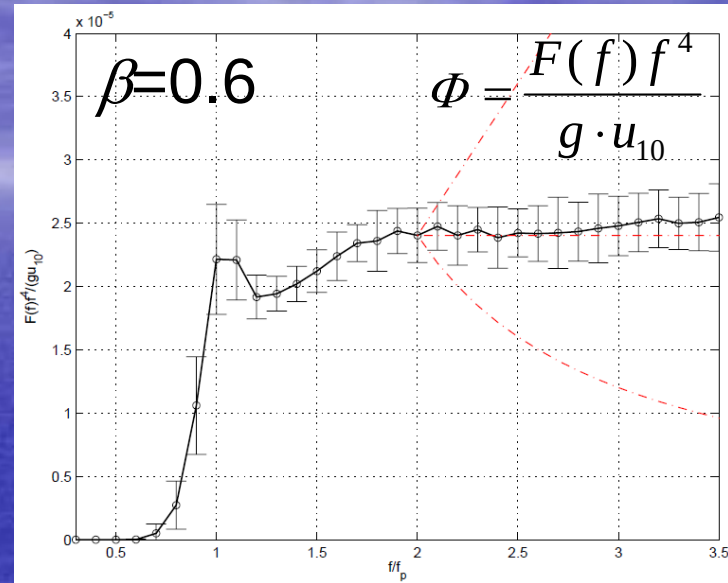
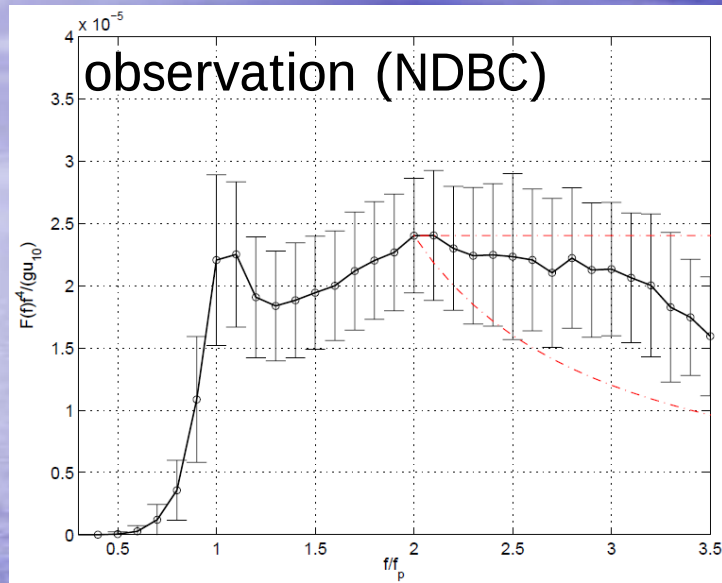
Saturation spectrum:

$$\Phi(f / f_p) = F(f) f^4 / gu_{10}$$

where $u_{10} / c_p > 1$

$$\beta=1$$

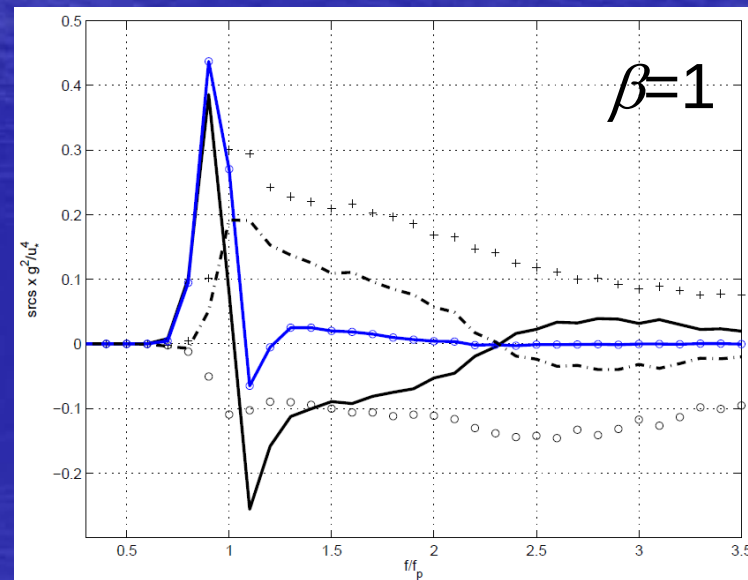
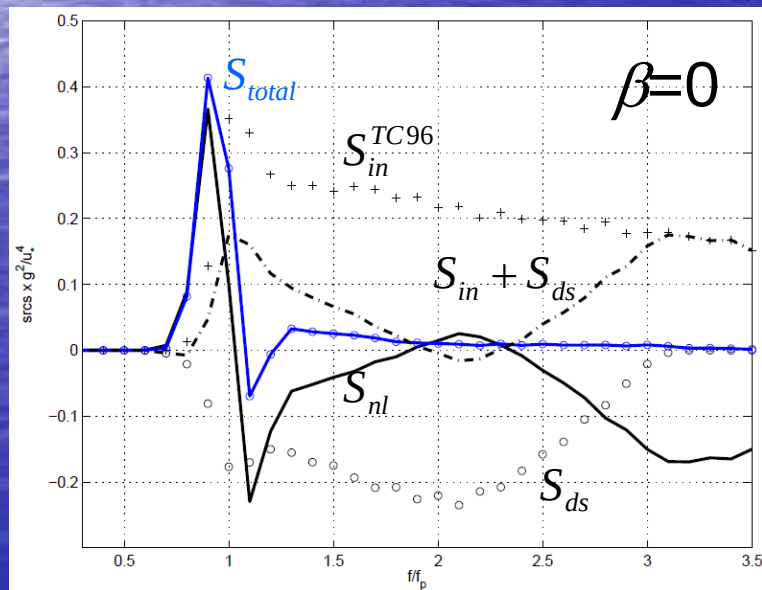
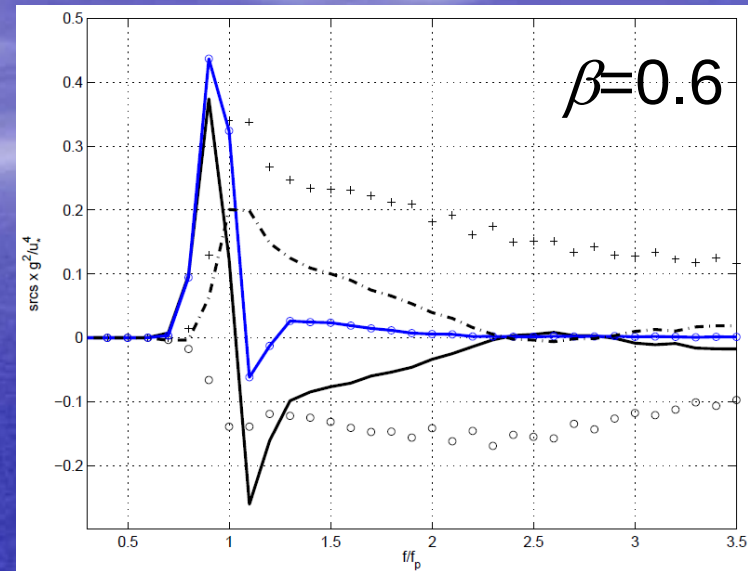
Wave spectrum in the equilibrium range



Source terms and their balance

$$S_{total} = S_{in}^{TC96} + S_{nl}^{SRIAM} + S_{ds}^{low} + \beta \cdot S_{ds}^{high}$$

normalized by $S^* = S \cdot g^2 u_*^{-4}$



Wave spectrum in the equilibrium range

Resio et al. (2004)

Any net gain or loss of energy within the equilibrium range would tend to force the spectrum away from an f^{-4} shape

Case 1

For a net gain ($S_{in} > S_{ds}$) and significant S_{nl} , the equilibrium spectrum would have to fall off less steeply than f^{-4}

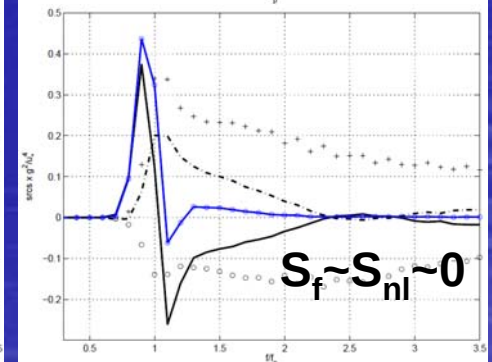
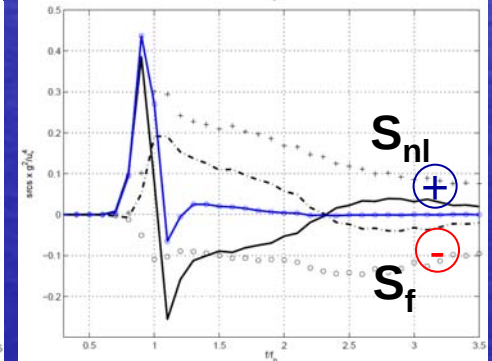
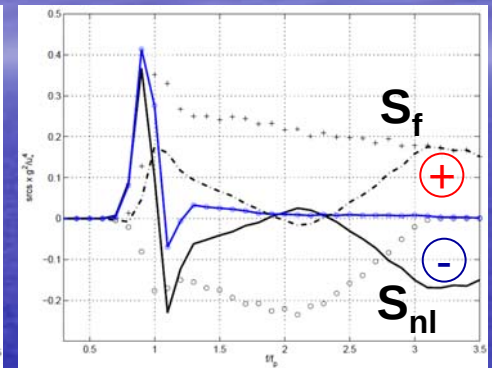
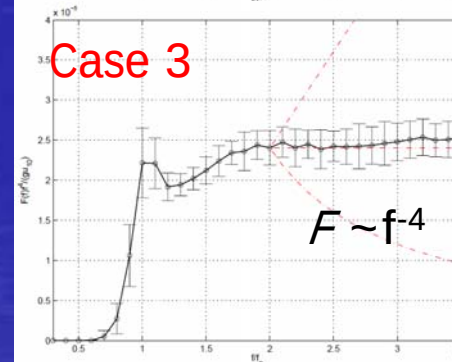
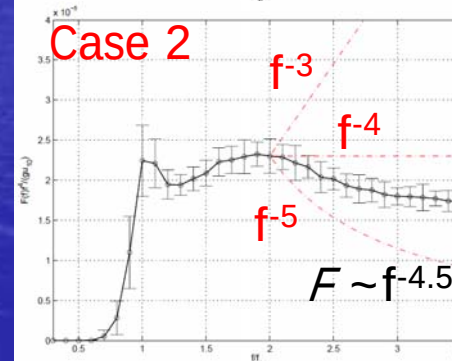
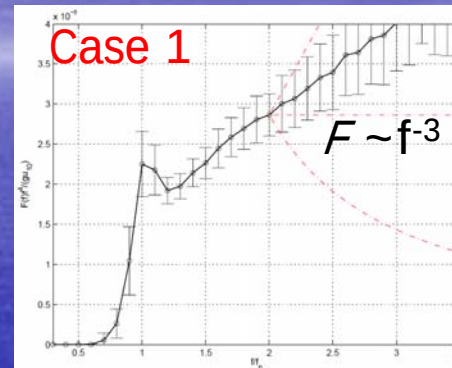
Case 2

For a net loss ($S_{in} < S_{ds}$) and significant S_{nl} , the equilibrium spectrum would have to fall off more steeply than f^{-4}

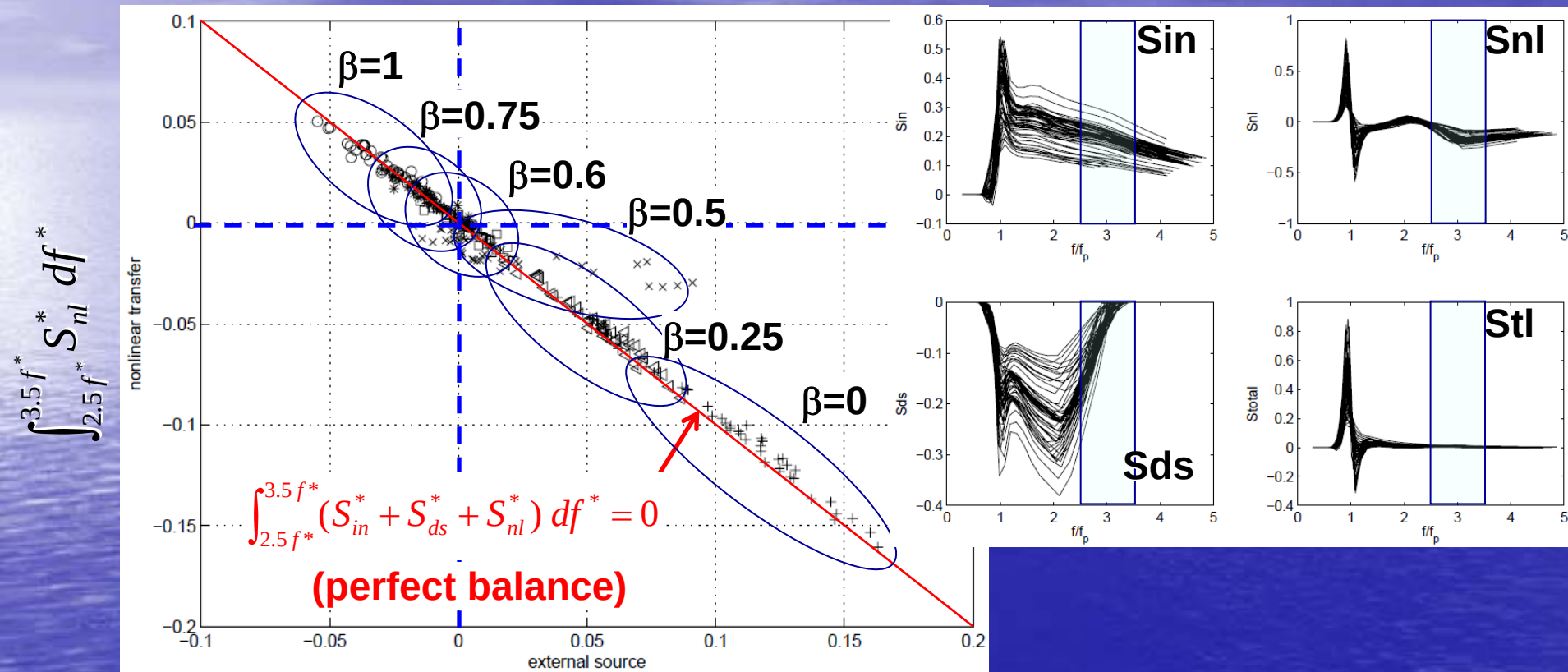
Case 3

If the net effect is negligible ($S_{in} + S_{ds} \sim 0$) within the equilibrium range, spectrum should be f^{-4} shape.

$$\int_{f_{eq}}^f \frac{\partial \xi^3}{\partial f} df = \xi^3(f) - \xi^3(f_{eq}) \sim \int_{f_{eq}}^f [S_{in} + S_{ds}] df$$



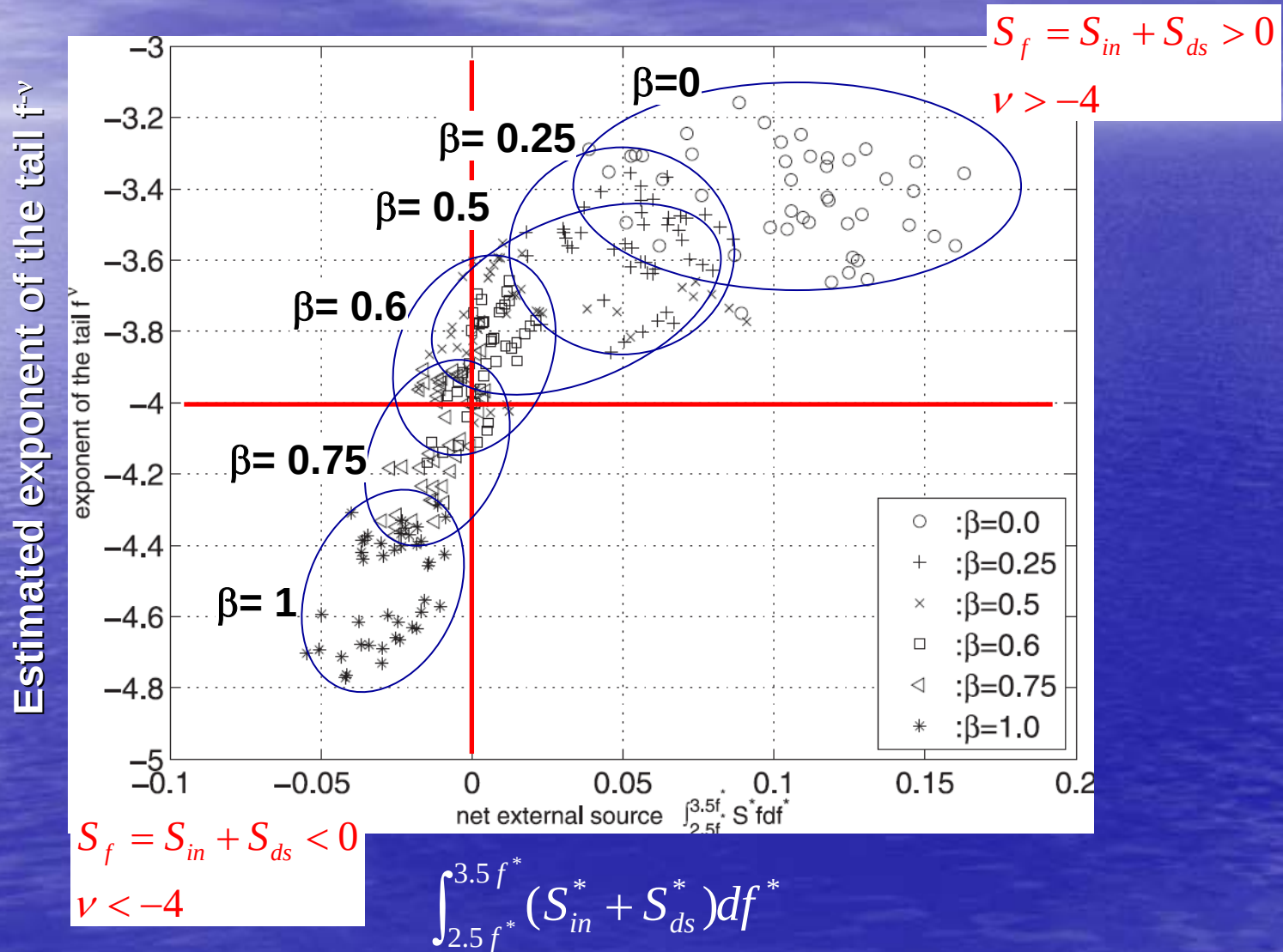
Intensity of the external source term and the S_{nl}



$$\int_{2.5f^*}^{3.5f^*} (S_{in}^* + S_{ds}^*) df^*$$

The total source term approached zero during the wave evolution whatever external source S_f is used.

External source term and the exponent of the spectral tail



Summary and discussions

Equilibrium condition: $S_{in} + S_{ds} + S_{nl} \approx 0$

Kitaigorodskii (1983)

$$S_{in} \sim S_{ds} \sim S_{nl} \sim 0$$

Phillips (1985)

$$S_{in} \sim S_{ds} \sim S_{nl}$$

Komen et al. (1984)

$$“S_{ds}” = -S_{in} - S_{nl}$$

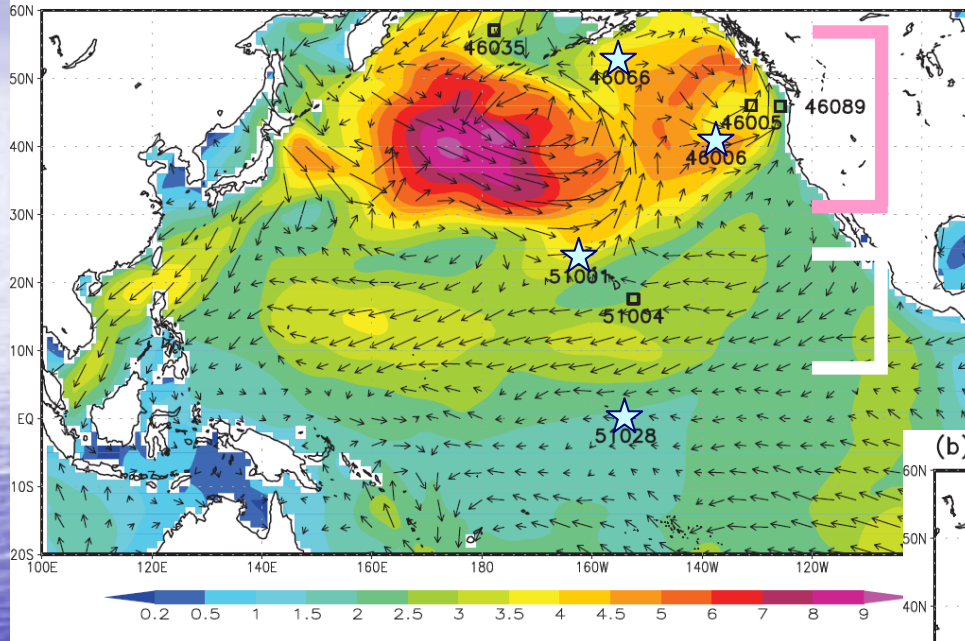
The sum of the three source terms approaches zero largely as a result of S_{nl} adjustment. However, the exponent of the spectral tail was also quite sensitive to S_f , in agreement with Resio et al (2004). The net external source S_f is the key factor that reproduces the f^{-4} tail.

Typical snapshot of Hs and Tp in the North Pacific

significant wave height

(a) Hs(m) and surface wind

2004/12/13 14z



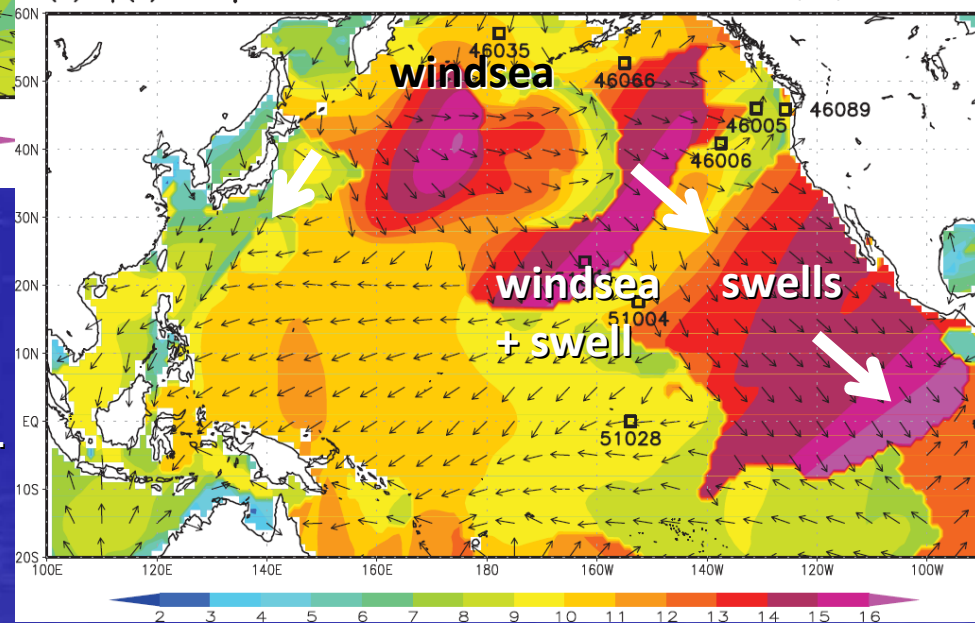
Enormous quantities of wind energy are transferred to surface waves in the mid-latitudes (associated with storm track)

Trade winds constantly generate local windsea in low latitudes

peak period

(b) Tp(s) and peak direction

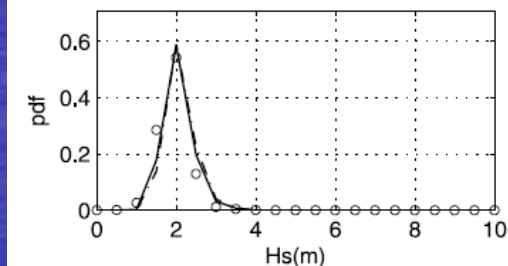
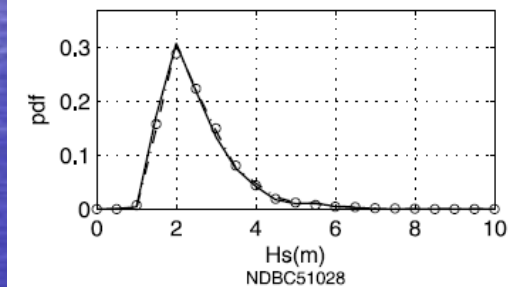
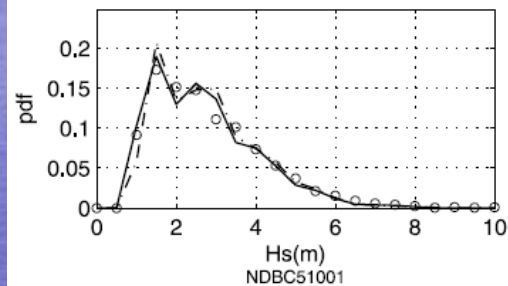
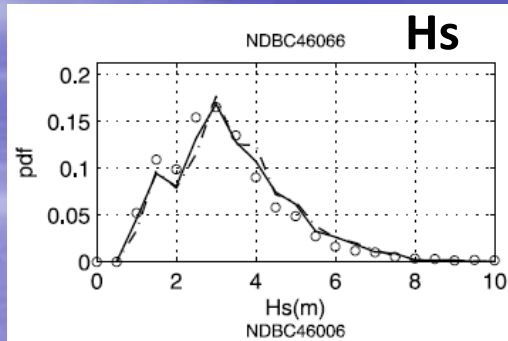
2004/12/13 14z



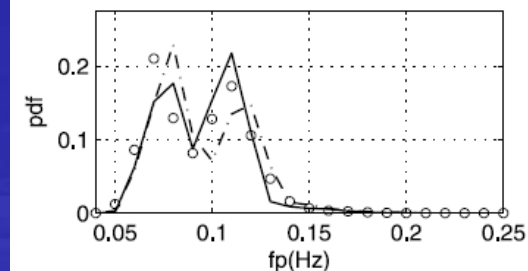
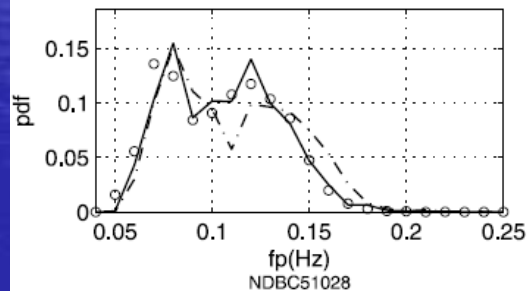
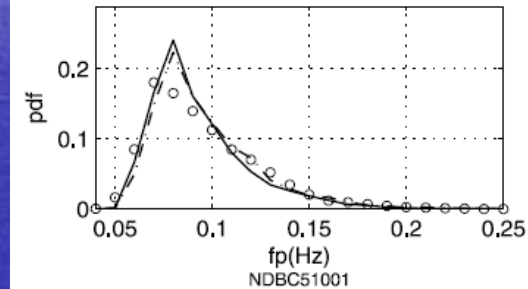
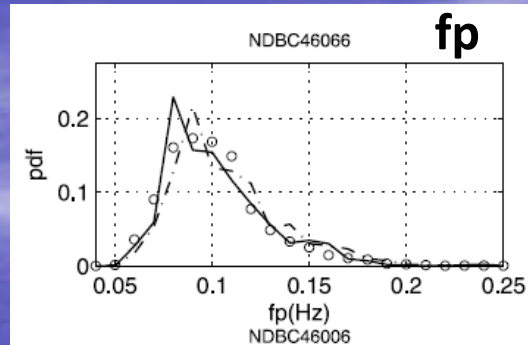
Ocean waves generated in the mid-lat. propagate to lower lat. as swells

The probability density functions of H_s and f_p

Higher
latitude



Lower
latitude



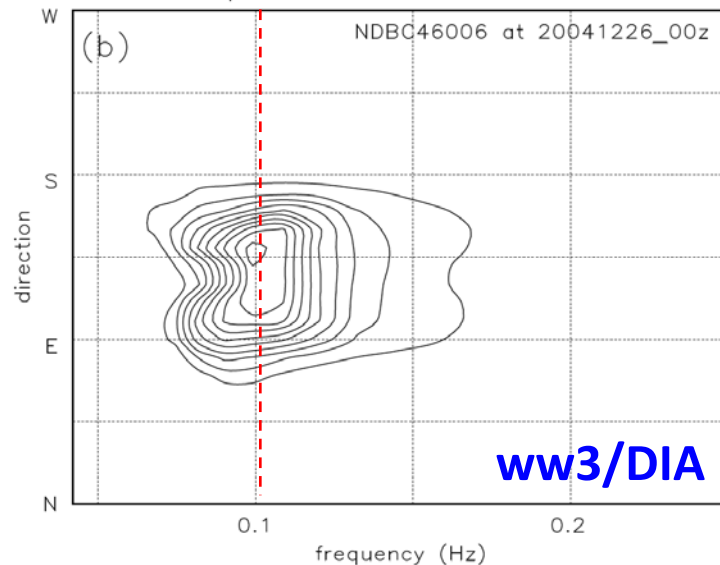
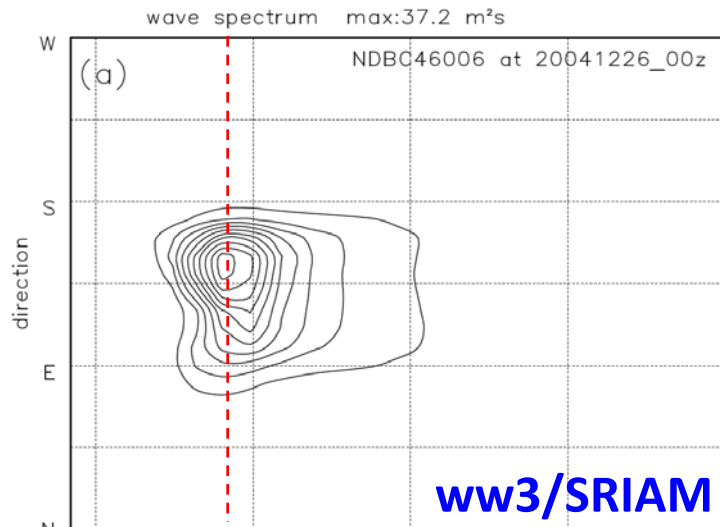
○ :in-situ data

— :ww3/SRIAM

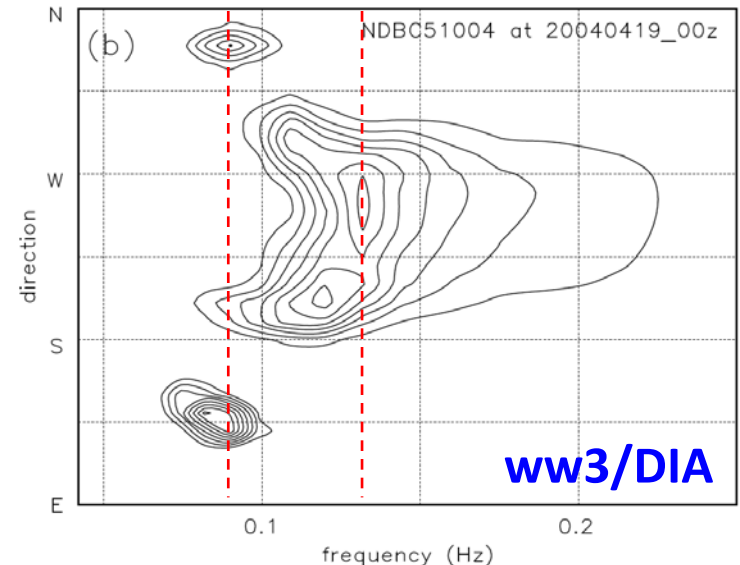
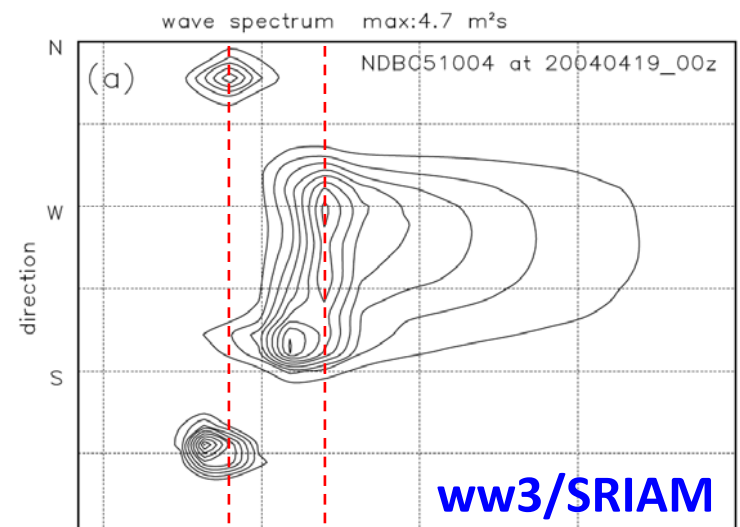
- - - :ww3/DIA

2D spectral shapes

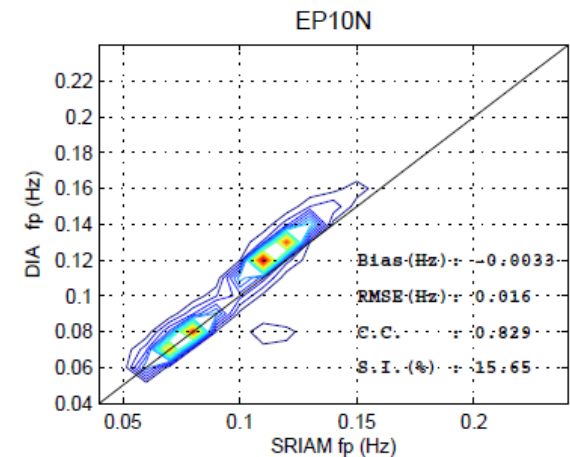
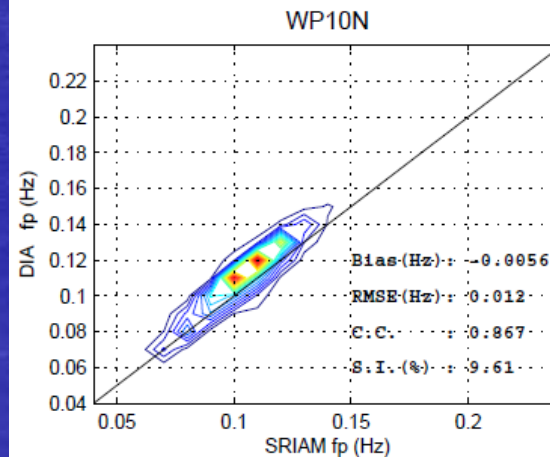
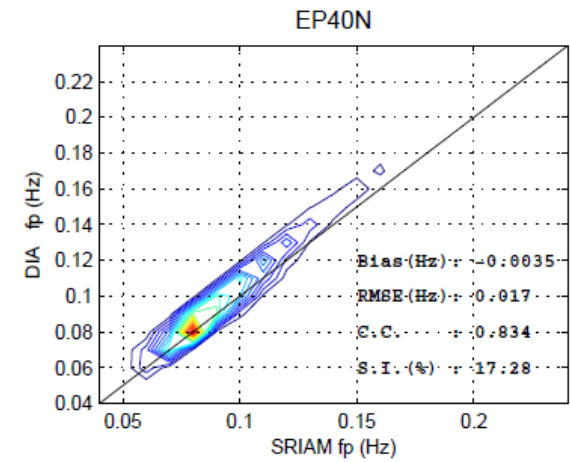
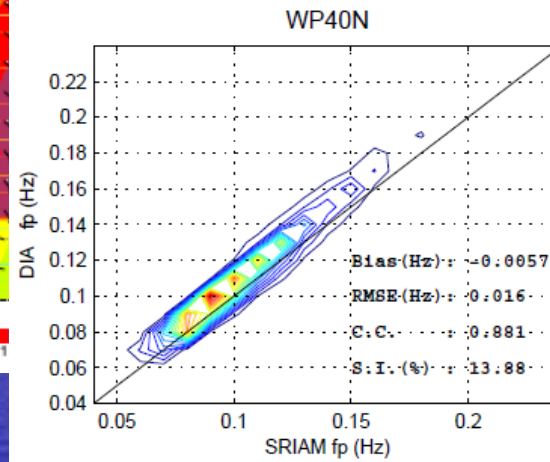
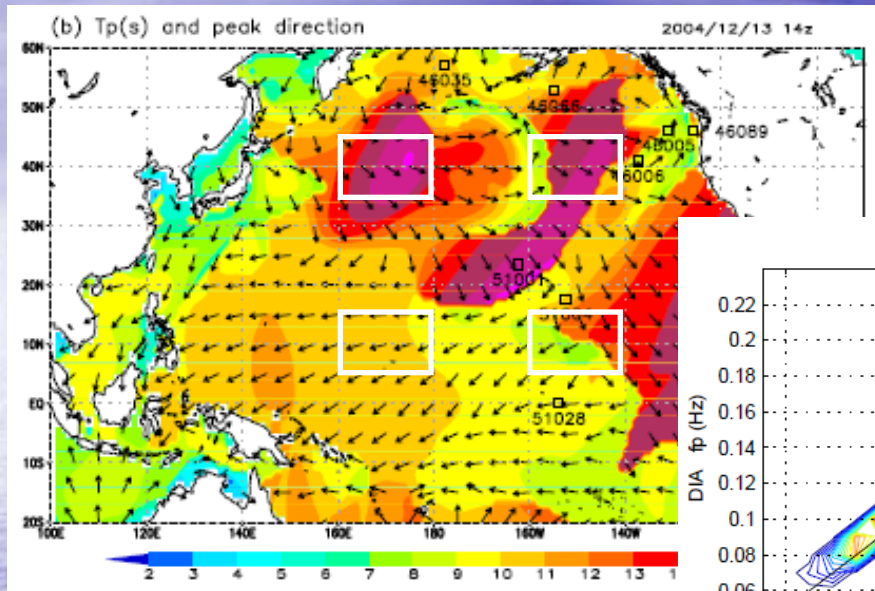
Station 46066 - S Aleutians 380NM



Station 51004 SE HAWAII 185 NM

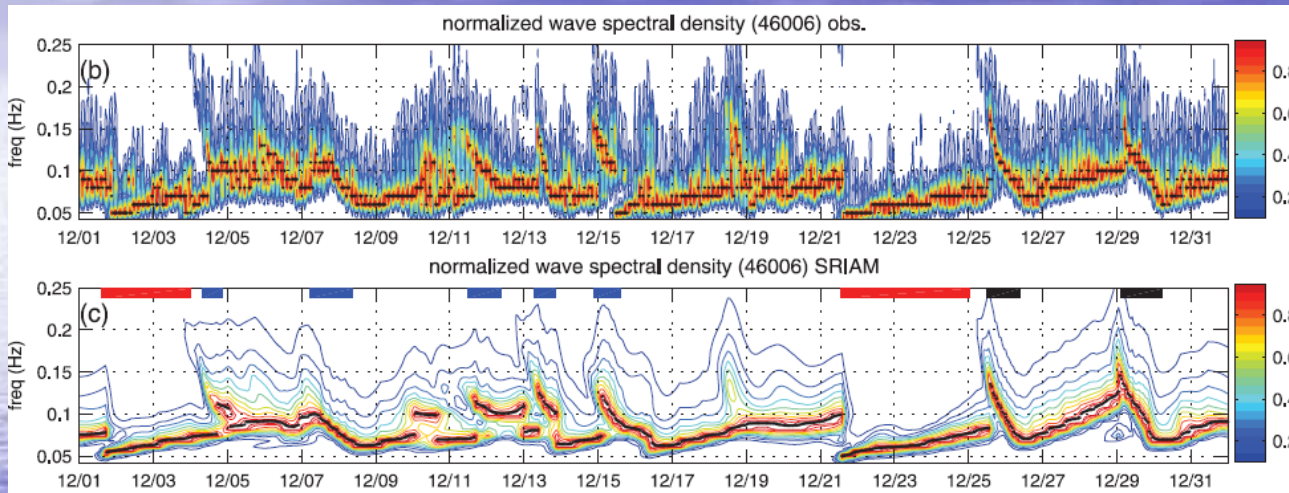


Joint PDFs of peak frequency



Evolution of wave spectra in-situ vs model (SRIAM)

in the mid-latitude Pacific (46006)



in the low-latitude Pacific (51004)

