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Interaction of Tsunamis with Short Surface Waves: An Experimental Study

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12th International Workshop on Wave
Hindcasting and Forecasting and 3rd Coastal
Hazard Symposium

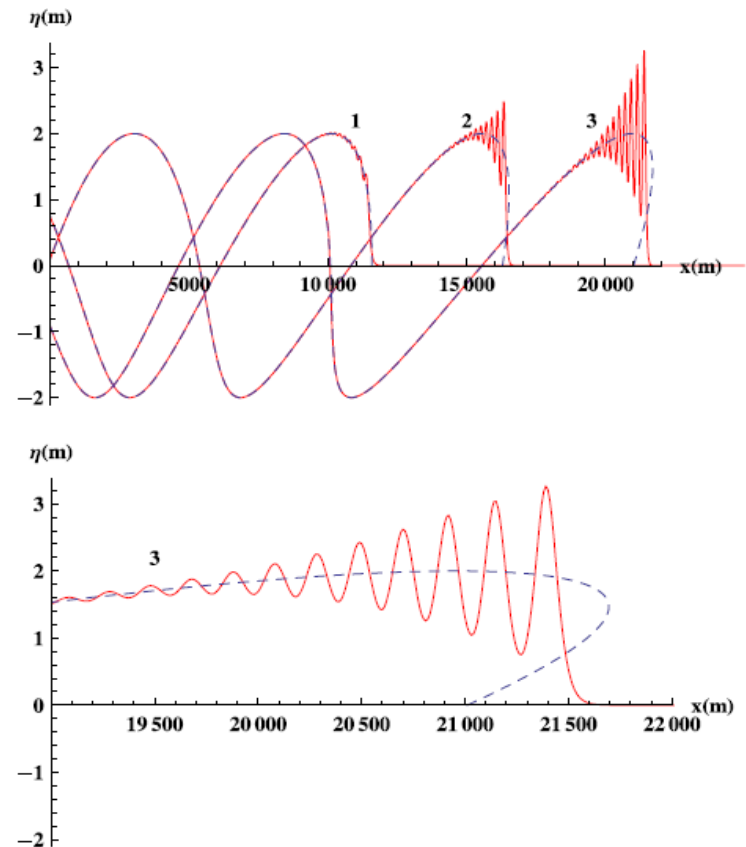
Motivation



2004 Indian Ocean Tsunami from Koh Jum Island

Motivation

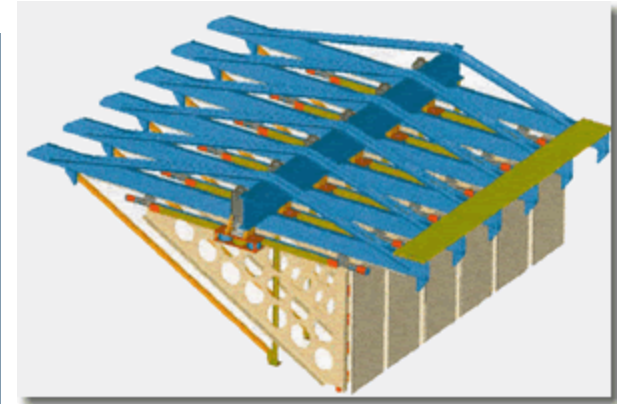
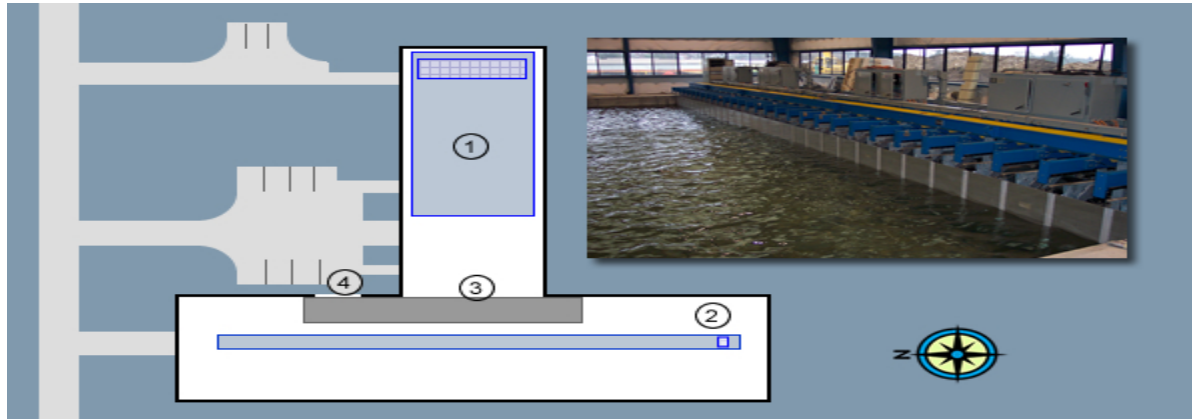
- Generally – default paradigm for tsunamis: **solitary wave**
 - Used for lab studies of tsunami damage
- Madsen et al. (2008)
 - Solitary wave paradigm flawed
 - Cause order of magnitude errors in spatial and temporal evolution over a sloping bottom



Interaction with the Swell Wave Field

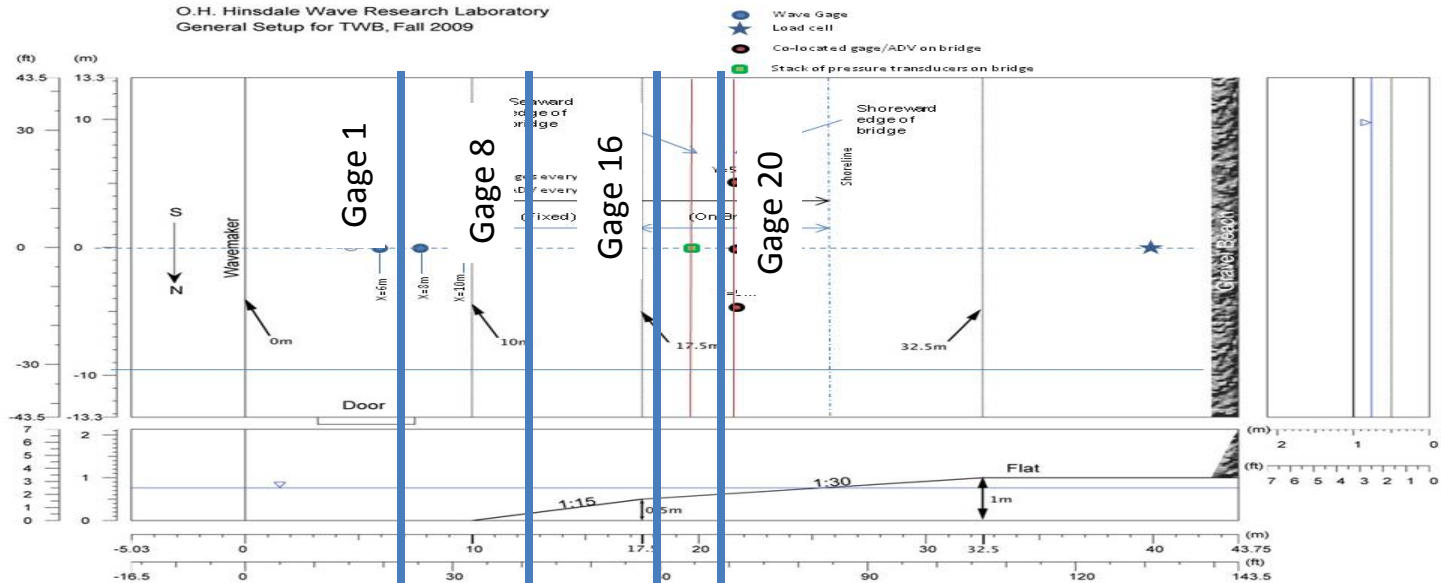
- Hypothesis:
 - Overlying swell wavefield can have some subtle effect on underlying tsunami
 - Front face steepness
 - Wave-wave interaction
- Run long wave through random swell field
- Use transient analysis (wavelet, Hilbert-Huang) to investigate effects

NEES Payload



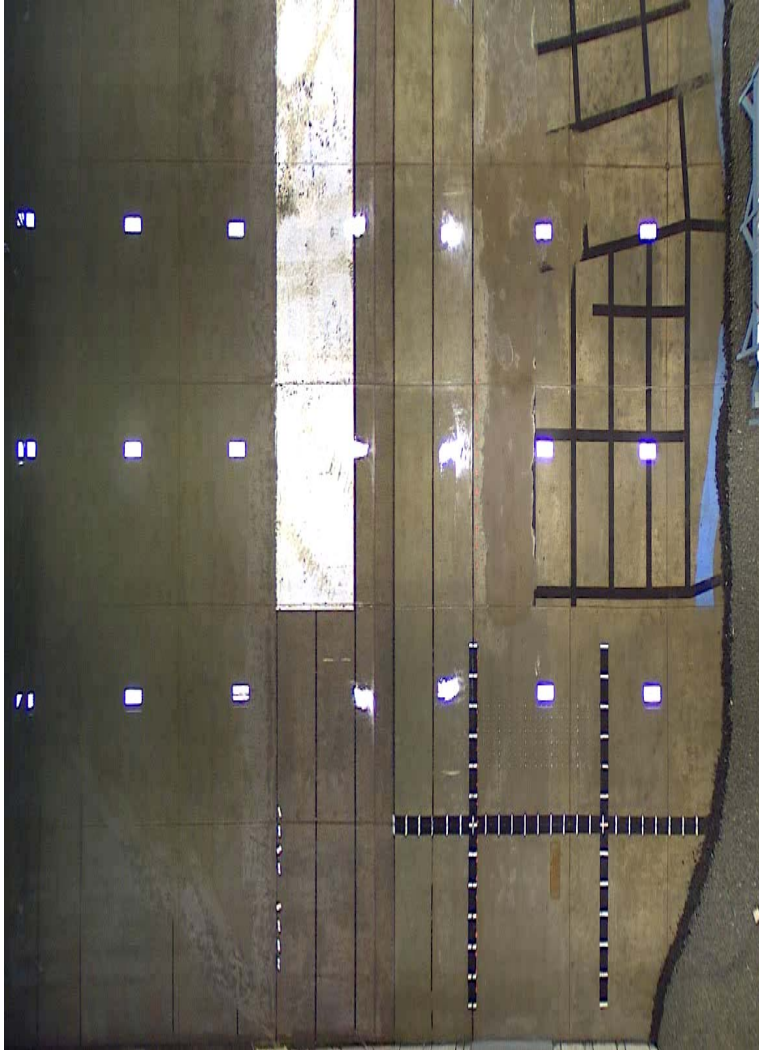
- One year project from NEES program, National Science Foundation
- Use NEES Tsunami facility at Oregon State University
- Tsunami Wave Basin:
 - 48.8m x 26.5m x 2.1m
 - 29-paddle multi-directional piston wavemaker
 - 4 resistance gages and 2 ADVs on movable bridge

Test Conditions

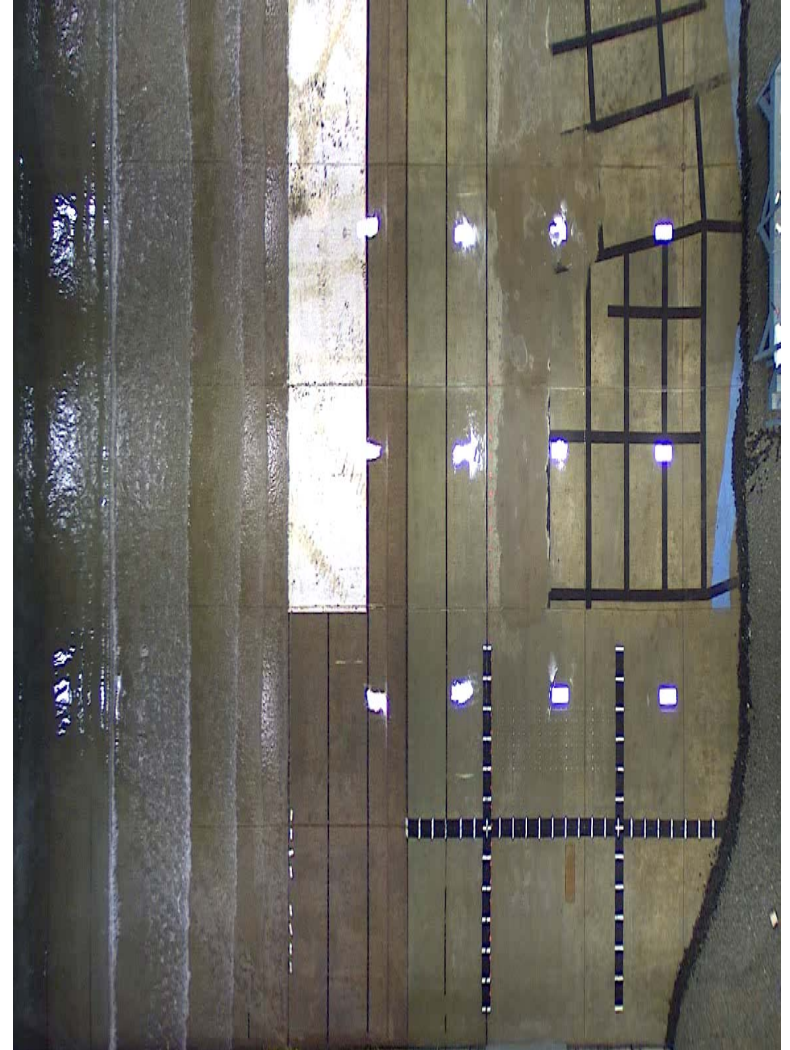


Test Number	Hs (cm)	Tp (s)	kh	$\delta=a/h$	U_r	Tsunami "height" ~30 cm Water depth 0.75 m Different runs with tsunami either in middle or end of swell record
1	5	2	1	0.033	0.033	
2	5	4	0.45	0.033	0.163	
3	10	2	1	0.067	0.067	
4	10	4	0.45	0.067	0.331	

Experiments



Tsunami Only

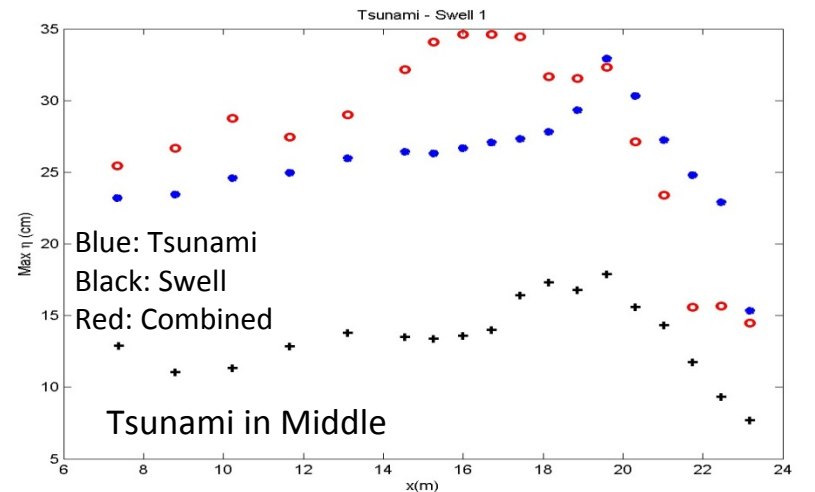
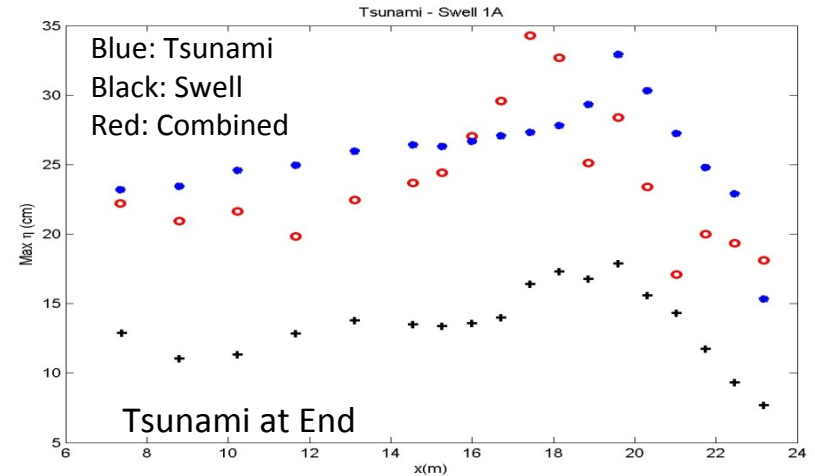
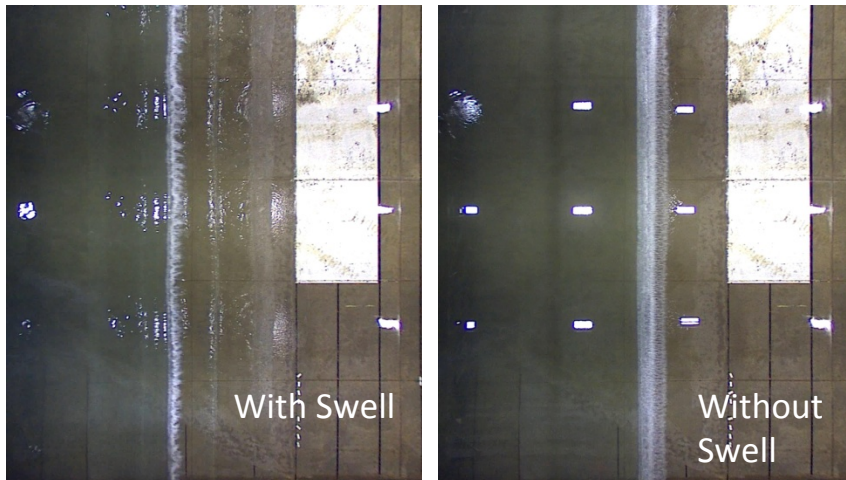


Tsunami with Swell

Analysis

Maximum Surface Elevation

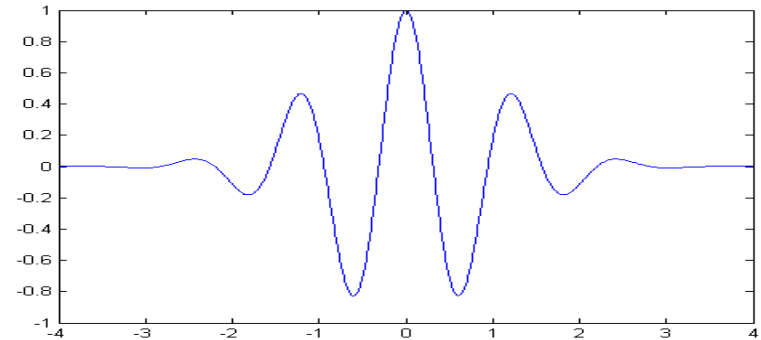
- As proxy for waveheight: maximum surface elevation in record
- Maximum surface elevation reached earlier with combined conditions than with either alone.
- Different results with tsunami in middle or at end of swell record.



Wavelet Analysis

Morlet Wavelet Transform

- Data from experiment – periodic signal interacting with a transient signal
- Standard Fourier Transform techniques not suitable
- Wavelet transform – time-dependent frequencies
- Spectral densities as a function of frequency and time

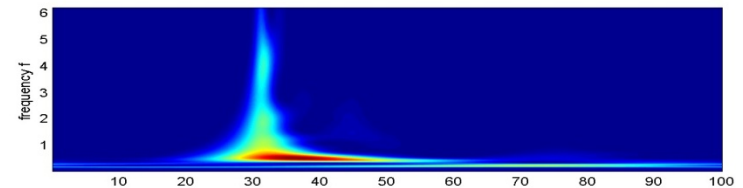
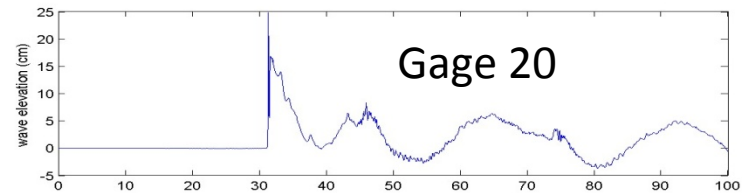
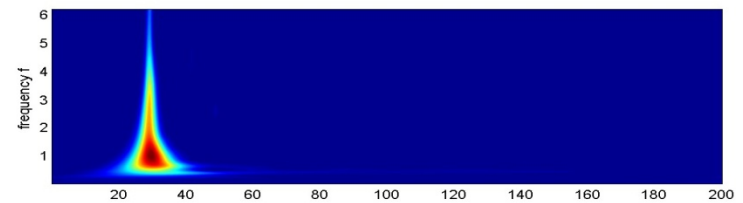
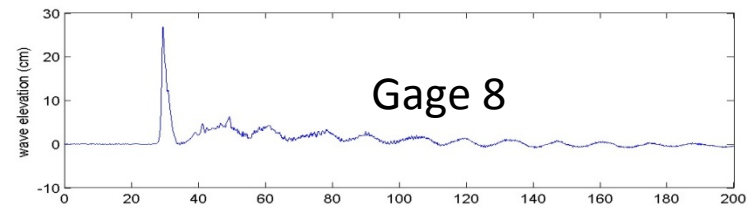
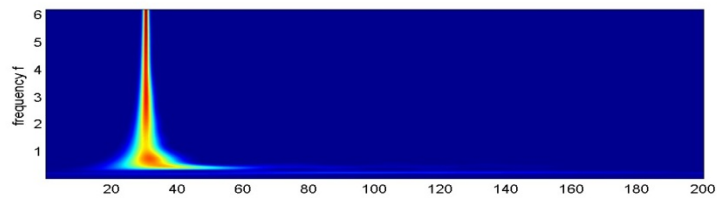
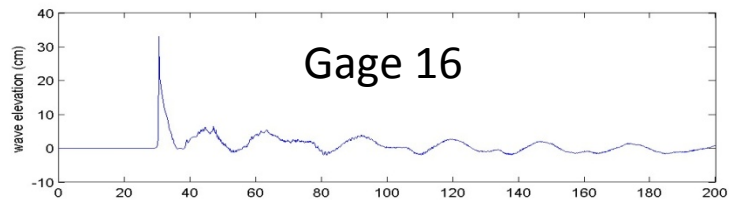
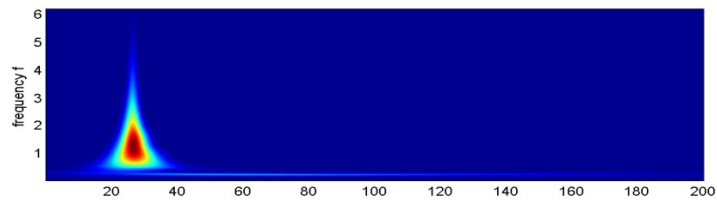
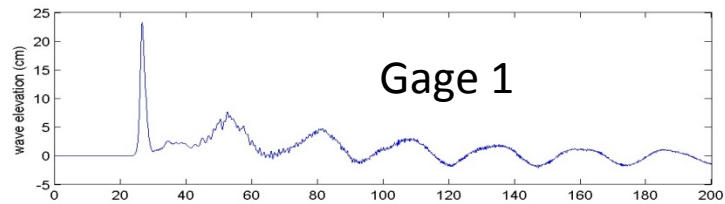


The Morlet Wavelet

$$\text{WT}(a, \tau) = \int_{-\infty}^{\infty} x(t) \psi_{a,\tau}^*(t) dt,$$

$$\psi(t) = \pi^{-1/4} \exp\left(-\frac{t^2}{2}\right) \exp(i\omega_0 t),$$

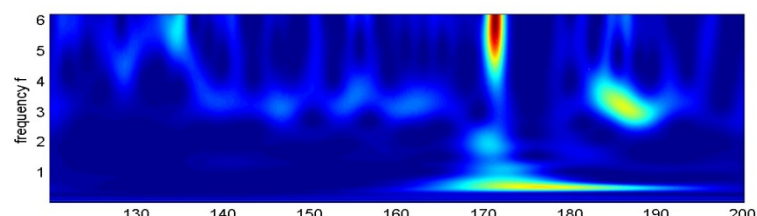
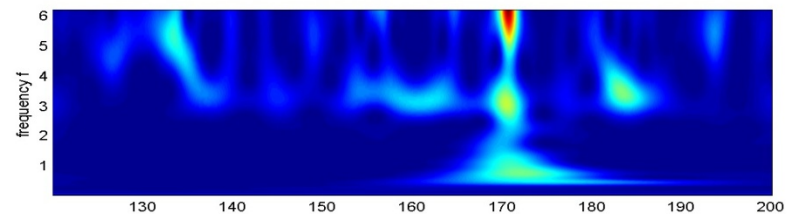
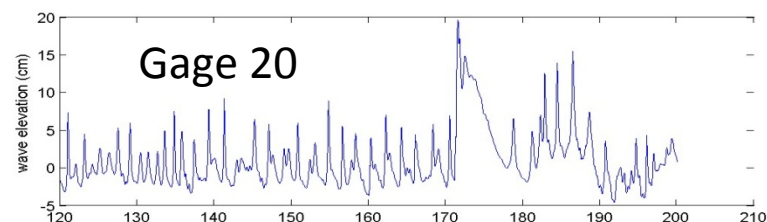
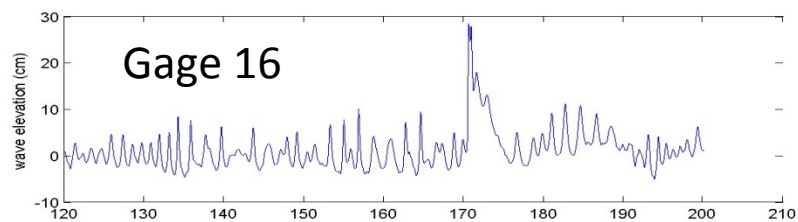
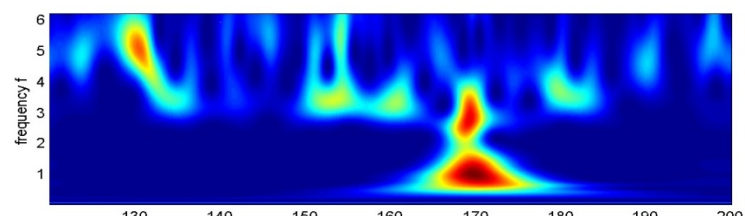
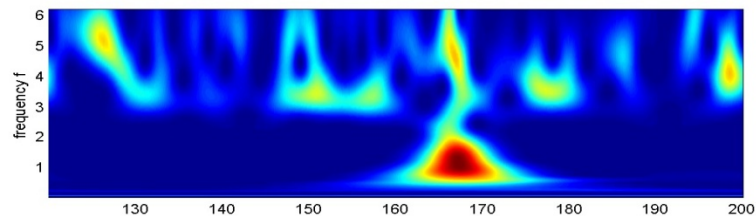
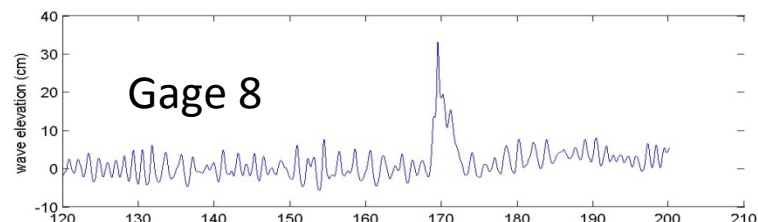
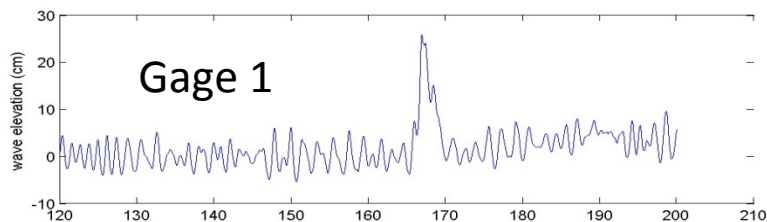
Wavelet Analysis



Tsunami
Only

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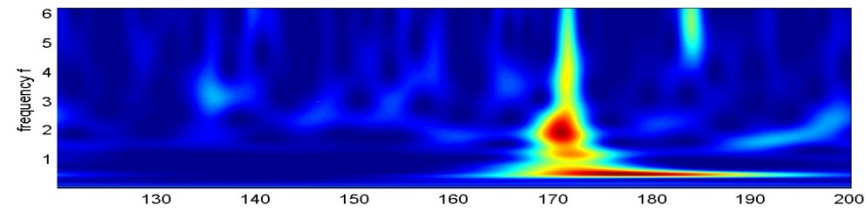
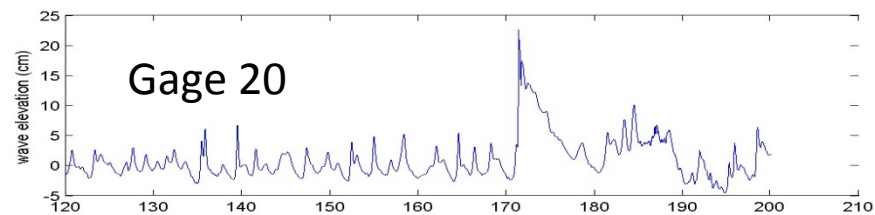
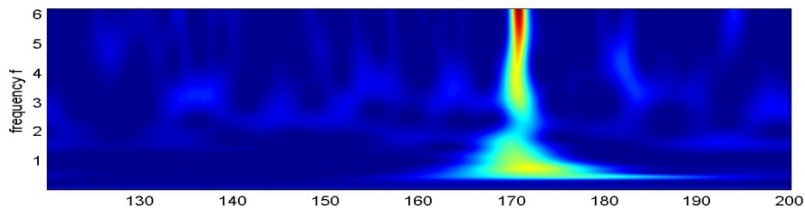
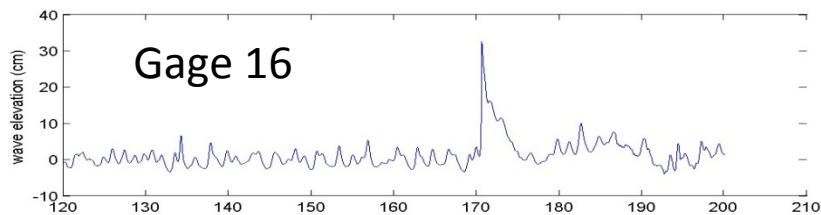
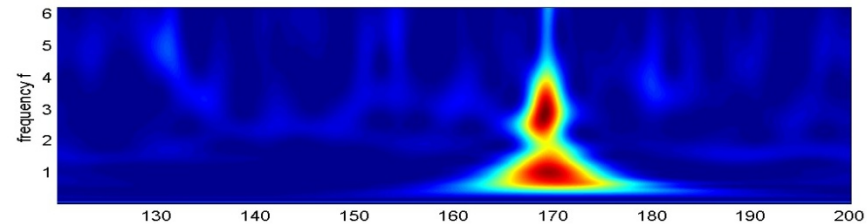
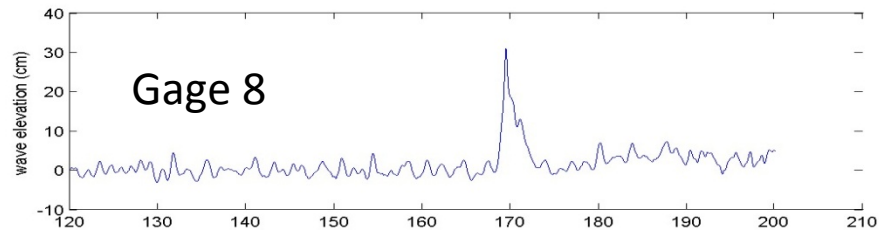
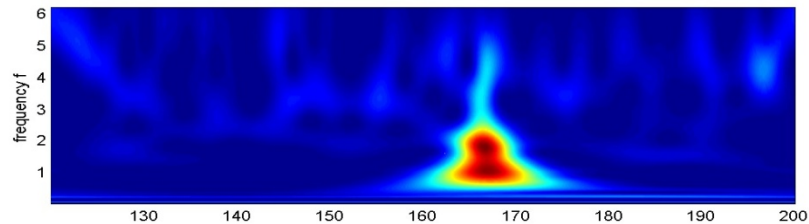
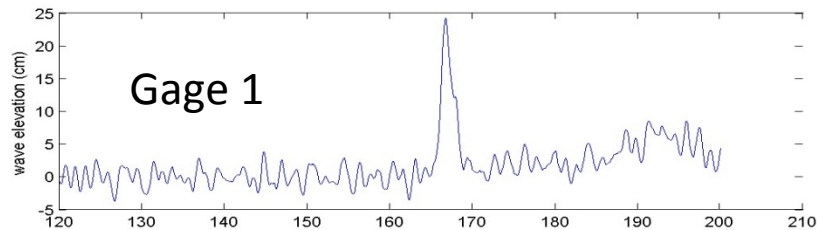
Wavelet Analysis



Tsunami with Swell: $H_s=5\text{cm}$,
 $T_p=2\text{s}$

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Wavelet Analysis



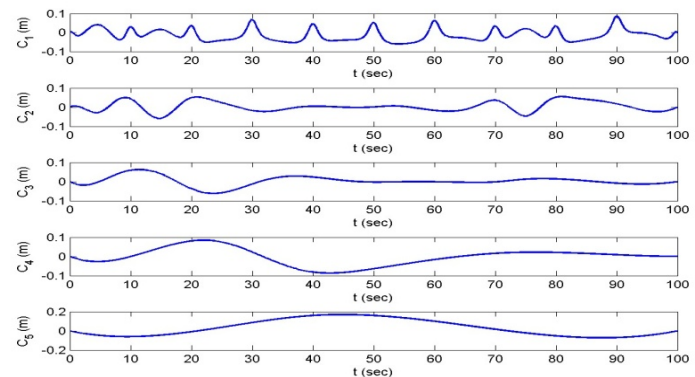
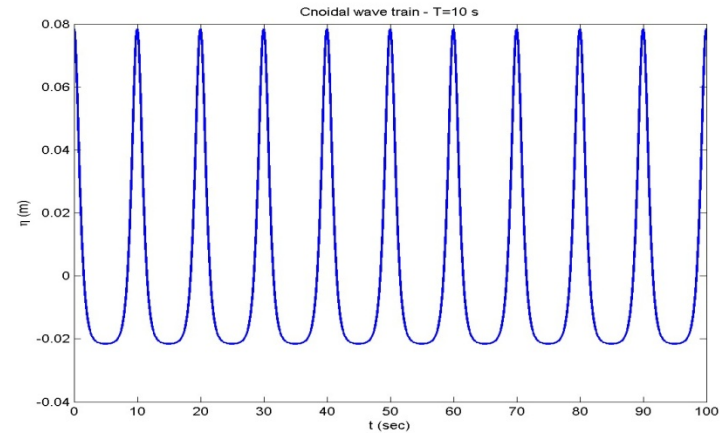
Tsunami with Swell: $H_s=5\text{cm}$,
 $T_p=4\text{s}$

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Hilbert-Huang Transform

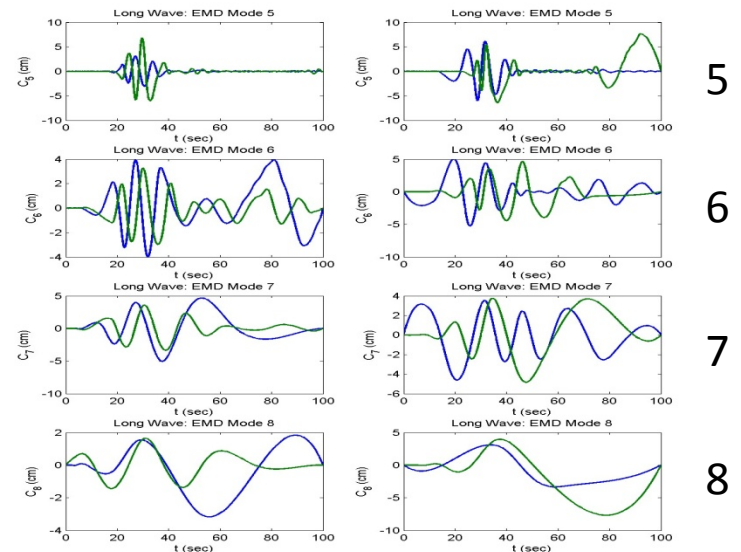
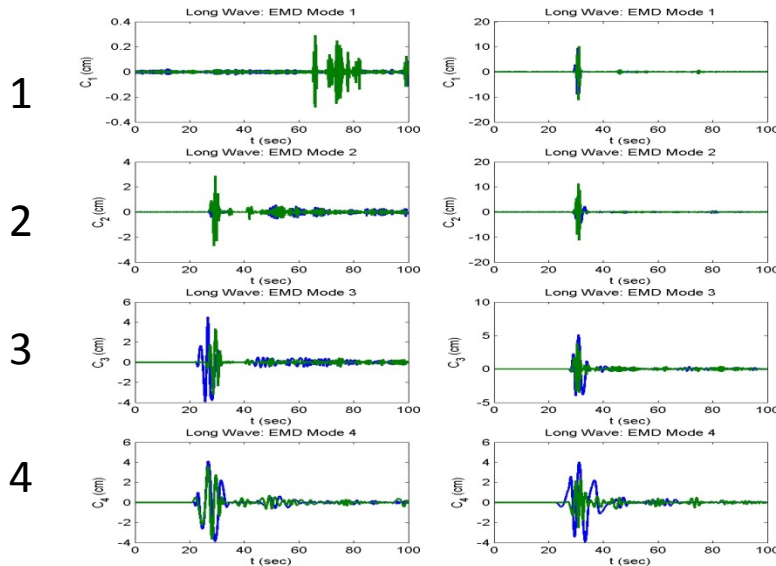
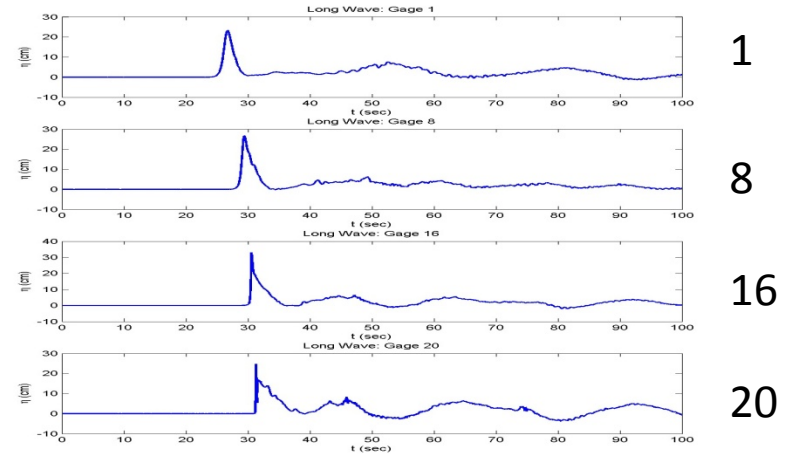
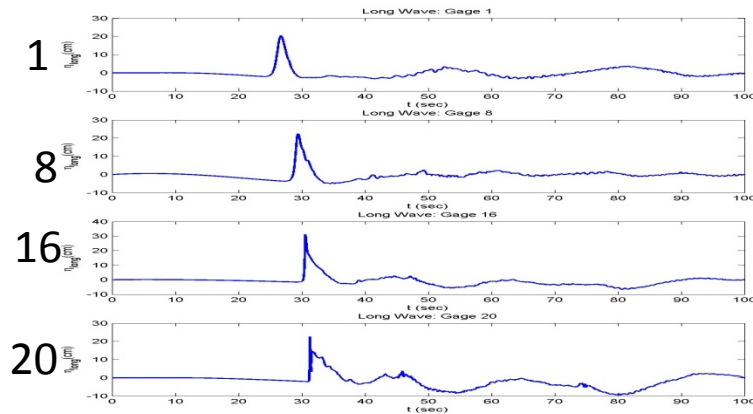
Hilbert Transform plus Empirical Model Decomposition

- Take original nonperiodic signal
- Calculate envelope with Hilbert Transform, then calculate mean
- Subtract mean from original signal
 - New signal
- Do this on each updated signal k times
- Result – mode c_1
- Subtract mode c_1 from original data – obtain new data
- Perform sifting process again
- Advantage:
 - No *a priori* basis function
 - Modes usually have some physical basis



Top: Cnoidal wave. Bottom. First five HHT modes of cnoidal wave.

HHT of Tsunami



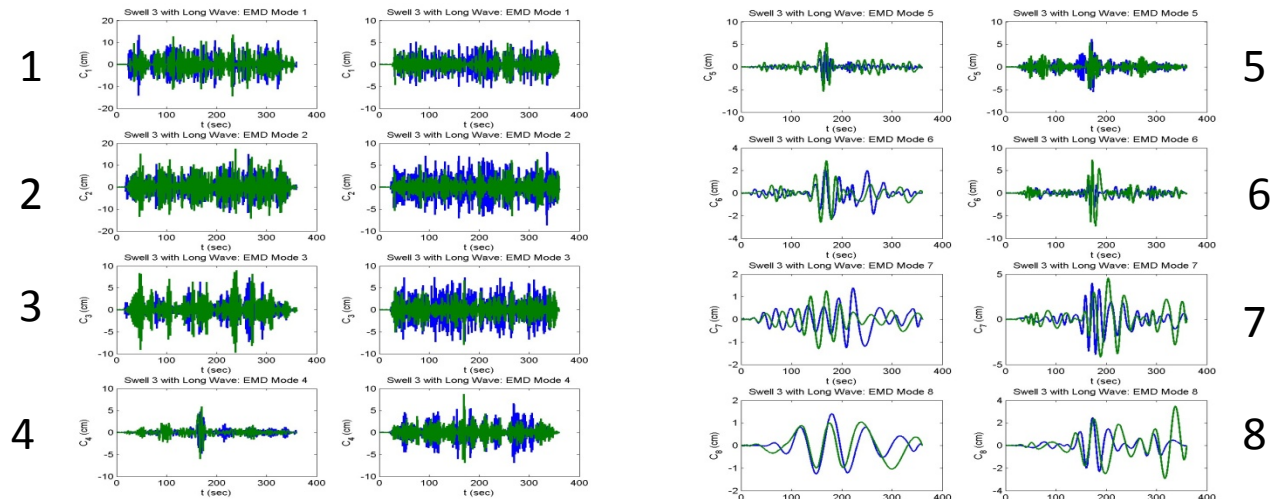
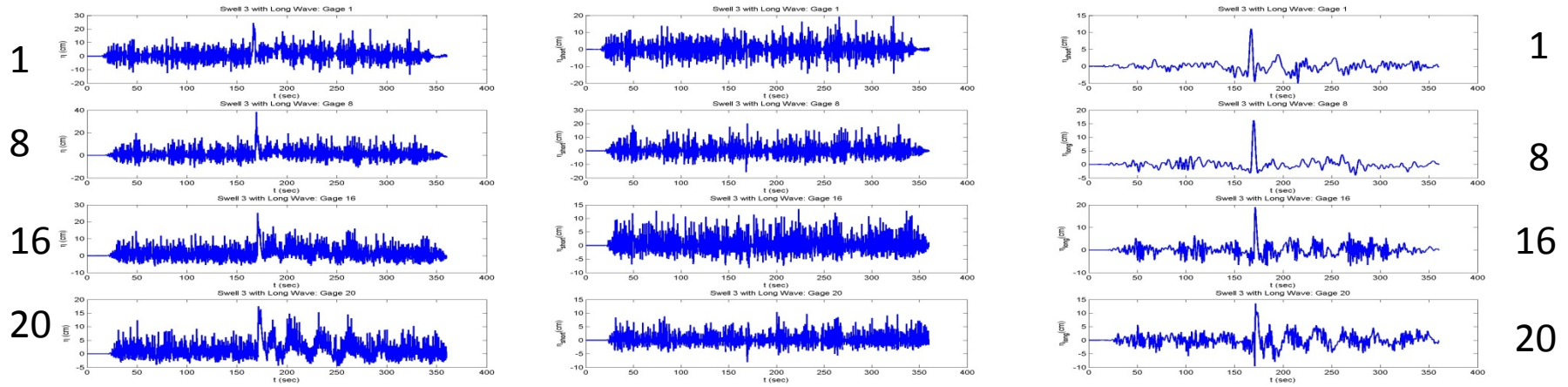
Blue: Gage 1
Green: Gage 8

Blue: Gage 16
Green: Gage 20

Blue: Gage 1
Green: Gage 8

Blue: Gage 16
Green: Gage 20

HHT of Combined Tsunami-Swell



Blue: Gage 1 Blue: Gage 16 Blue: Gage 1 Blue: Gage 16
 Green: Gage 8 Green: Gage 20 Green: Gage 8 Green: Gage 20

Steepness

Breaking: related to wave steepness $\frac{\partial \eta}{\partial x}$:

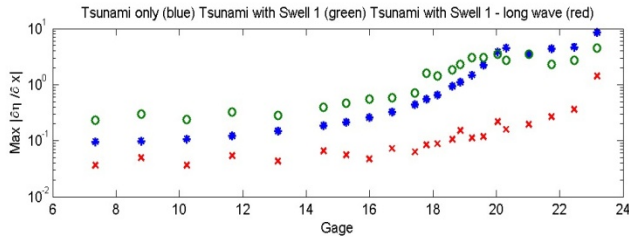
- Deduce steepness from:
 - Time series
 - Combined HHT modes (long/short motions)
- Look for maximum steepness

Approximate
steepness by linear
wave equation:

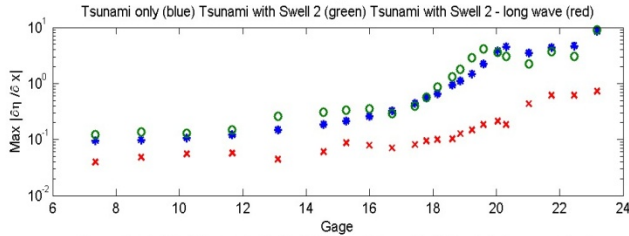
$$\frac{\partial \eta}{\partial x} = -\frac{1}{\sqrt{gh}} \frac{\partial \eta}{\partial t}$$

Steepness

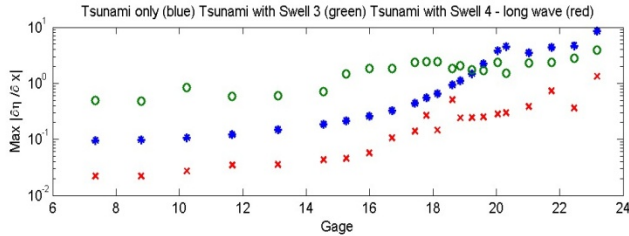
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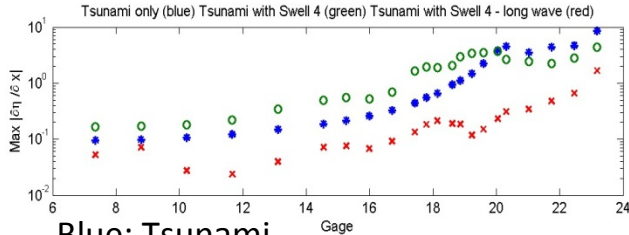
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3



4

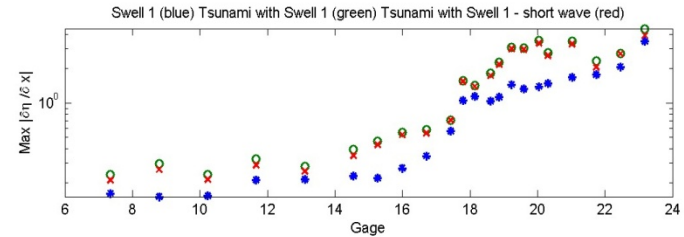


Blue: Tsunami

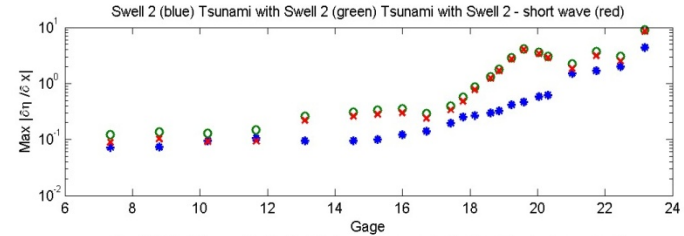
Green: Tsunami with swell

Red: Tsunami with swell – long wave

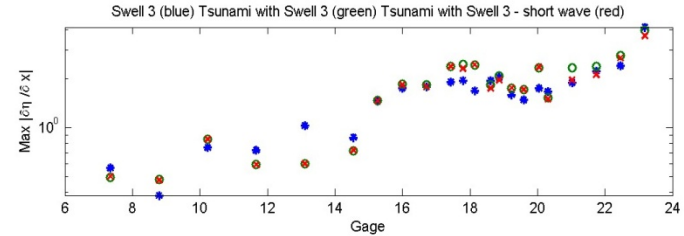
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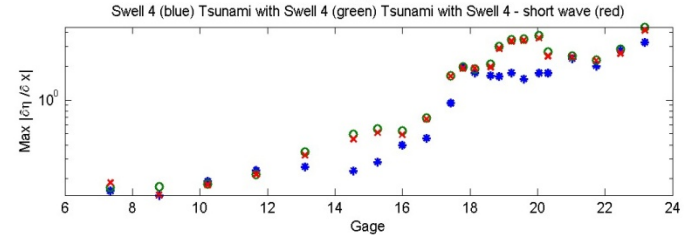
2



3



4



Blue: Swell only

Green: Swell with tsunami

Red: Swell with tsunami – short wave

Conclusions Thus Far

- It was hypothesized that the overlying swell field might have some effect on a long tsunami (front face steepness, etc.)
- Experiments were run at the NEES Tsunami Facility at Oregon State University
- Some indication of swell affecting maximum surface amplitude of overall wave field
- Wavelet analysis: spectral structure of tsunami affected by swell
- Use Hilbert-Huang Transform to break motion up into scales
- Analysis of steepness in combined motion: strong energy shift to high frequencies made more efficient with swell present.